

Impacts of anthropogenic and natural NO_x sources over the U.S. on tropospheric chemistry

Renyi Zhang*[†], Xuexi Tie[‡], and Donald W. Bond*

*Department of Atmospheric Sciences, Texas A&M University, College Station, TX 77843; and [‡]Atmospheric Chemistry Division, National Center for Atmospheric Research, Boulder, CO 80307

Communicated by Mario J. Molina, Massachusetts Institute of Technology, Cambridge, MA, December 12, 2002 (received for review August 9, 2002)

We evaluate the impact of anthropogenic and natural NO_x sources over the contiguous United States on tropospheric NO_x and O₃ levels by using a global 3D chemical transport model. The effects of major U.S. surface NO_x emission sources (including anthropogenic, biomass burning, and soil emissions) are compared with that of lightning-produced NO_x. Summer lightning is shown to play a dominant role in controlling NO_x and O₃ concentrations in the middle and upper troposphere, despite the fact that fossil-fuel burning represents the largest source of NO_x over the U.S. Furthermore, the effect of regional U.S. lightning is propagated through large areas of the Northern Hemisphere by atmospheric circulation. The results reveal that a thorough assessment of atmospheric NO_x emission sources and their impact is required to devise control strategies for regional and global air pollution.

Nitrogen oxides (NO_x = NO + NO₂) are atmospheric catalysts that are related to ozone and hydroxyl radicals (1–3). The abundance of NO_x regulates the atmospheric oxidizing power and global biogeochemical cycles (4, 5). In the troposphere, NO_x is closely related to the ozone chemistry via two separate processes. In regions of high NO_x concentrations ozone is produced photochemically in the cycling of NO to NO₂, which is facilitated by peroxy radicals formed during oxidation of carbon monoxide, methane, and volatile organic compounds, whereas in regions of low NO_x concentrations ozone is catalytically destroyed. Nitrogen oxides are also intricately linked to the hydroxyl radical OH, another key atmospheric oxidizing species. The reaction between NO₂ and OH leads to the formation of relatively stable nitric acid HNO₃, which can be removed from the atmosphere by precipitation and provides fixed nitrogen for the biosphere. Also, because O₃ absorbs strongly the Earth's IR radiation, knowledge of the regional and global NO_x distribution is important for climate studies (6).

NO_x is emitted into the atmosphere from natural and anthropogenic sources, i.e., from fossil fuel combustion, biomass burning, oxidation of atmospheric ammonia, and lightning (3, 7). Transport of NO_x from the stratosphere and aircraft emissions could also be important sources in the upper troposphere (8, 9). Over the contiguous U.S., anthropogenic emissions from fossil-fuel burning of automobiles, power plants, or industries constitute the main source of NO_x, with an estimated production of 7.3 Tg N yr⁻¹ (1 Tg = 10¹² g) (10, 11). Other NO_x emission sources over the U.S. include biomass burning and soil release, which are about 0.3 and 0.5 Tg N yr⁻¹, respectively (10, 11). Lightning accounts for ≈5% of the total U.S. NO_x emissions (0.4 Tg N yr⁻¹) and the lightning emission increases in the summer, reaching ≈14% of the total emission in July (7). Compared with the global NO_x emission of 32 Tg N yr⁻¹ (with 21 Tg N yr⁻¹ attributed to fossil-fuel burning) (10, 11), the U.S. contributes to a significant fraction of the world's total anthropogenic NO_x production.

In this study we assess the impacts of anthropogenic and natural NO_x sources over the continental United States on tropospheric chemistry by using a 3D global numerical model, Model of Ozone and Related Chemical Tracers (MOZART, version 2) (12). The tropospheric budgets of NO_x and O₃ are

analyzed to determine the relative contributions of the various NO_x sources over the U.S.

Methods

The meteorological data for MOZART is produced from the output of the National Center for Atmospheric Research community climate model (12). MOZART accounts for advection, convection, cloud formation, diffusive exchanges, gas phase and heterogeneous chemistry, photochemical transformations, wet deposition, and dry deposition processes. MOZART includes biogenic emissions of chemicals such as isoprene, carbon monoxide emissions from the ocean surface, and aircraft emissions of NO_x. The global and U.S. NO_x surface emissions (i.e., from anthropogenic activity, biomass burning, and soil release) are taken from recent compilations by Olivier and colleagues (10, 11). NO_x production by lightning over the U.S. is based on ground- and satellite-based lightning measurements (7). Monthly cloud-to-ground (CG) flashes are obtained by averaging the National Lightning Detection Network (NLDN) data over the 5-year period 1995–1999, and intracloud (IC) flashes are obtained from the observed geographical distribution of the IC/CG ratio based on NLDN and satellite optical transient detector (OTD) measurements. The IC and CG flashes are distinguished according to the procedure described in refs. 7, 13, and 14. IC flashes are believed to have significantly lower energies than CG flashes and hence have a lower NO_x production efficiency (15). Lightning production outside the U.S. is based on parameterization developed by Price *et al.* (15) and satellite observation from OTD (13). The annual production of NO_x by lightning used in this study are 0.4 and 7.0 Tg N yr⁻¹ over the U.S. and the whole globe, respectively. Lightning-produced NO_x is distributed uniformly within the vertical extent of the convective clouds. The distributions of lightning-produced NO_x are slightly extrapolated to coincide with the modeled convective cloud distributions. The uncertainty of NO_x production by lightning has been evaluated in our previous studies (7).

Results and Discussion

To assess the importance of the various NO_x sources over the U.S. on tropospheric chemistry, we simulate the concentrations of NO_x and O₃ by using MOZART. Three model runs are considered: (i) with both lightning and surface (including anthropogenic, biomass burning, and soil release) emission sources of NO_x, (ii) without the lightning source, and (iii) without the surface emission sources within the boundaries of the continental U.S. Fig. 1 illustrates the zonally averaged percentage changes in NO_x and O₃ concentrations in July calculated because of lightning or surface emission sources. The lightning impact is determined by subtracting the concentrations derived by model run ii from those of model run i (Fig. 1 *Upper*). Similarly, the surface emission impact is evaluated by using the results from the difference of model

Abbreviations: CG, cloud-to-ground; IC, intracloud.

[†]To whom correspondence should be addressed. E-mail: zhang@ariel.met.tamu.edu.

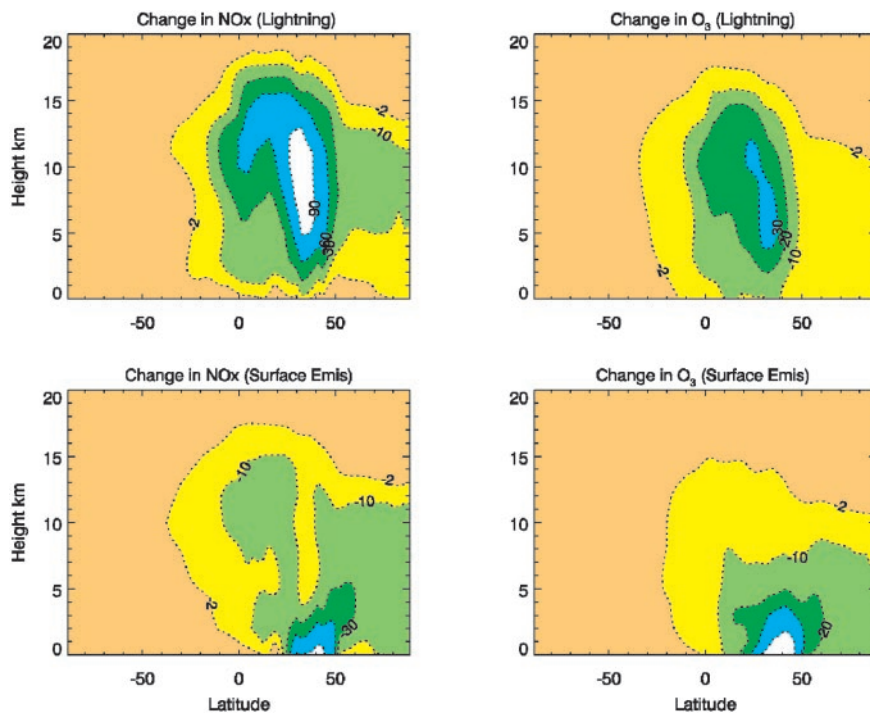


Fig. 1. Zonally averaged percentage changes in NO_x and O₃ concentrations caused by lightning (*Upper*) and surface (*Lower*) NO_x emissions (see text). Results are averaged over the month of July and the longitudes correspond to the continental U.S. (75–125° W).

runs *iii* and *i* (Fig. 1 *Lower*). Lightning has a significant impact on the concentration of NO_x in the middle and upper troposphere. Between 5 and 15 km inclusion of lightning NO_x production leads to a change of >90% in the NO_x concentration between 30° and 40° N. The model results also indicate that the impact of U.S. lightning is propagated throughout the Northern Hemisphere and into the Southern Hemisphere. A change of >60% in the NO_x concentration occurs at 10–12 km near the equator, and a change of >30% extends as far as 10° S at 12 km. On the other hand, the impact of the U.S. surface NO_x sources is important only in the lower troposphere (<5 km). In July lightning also strongly affects upper tropospheric O₃, resulting in a change of 30% or more in the ozone concentration from 4 to 12 km within 30–40° N. An area of >20% changes extends across the equator at 10 km. As is the case with NO_x, the influence of surface NO_x sources on O₃ is limited primarily at lower levels.

To quantify the relative importance of lightning and surface NO_x emissions over the U.S., monthly budget analyses are performed by computing the column mass-weighted percentage changes in concentrations of NO_x and O₃, depicted in Fig. 2. The summations are calculated for the latitudinal and longitudinal domain of the continental U.S. over three altitude ranges: 0–5 km (lower troposphere), 5–10 km (upper troposphere), and 0–10 km (entire troposphere). In the lower troposphere, contribution of lightning to the NO_x budget peaks at ≈20% in the summer months and diminishes in the winter; throughout the year anthropogenic emissions dominate the NO_x budget. In the upper troposphere, however, the lightning impact is significant in the spring, summer, and fall, i.e., >50% from April to November and ≈90% in July and August. In July and August, the entire tropospheric NO_x budget is nearly equally contributed by lightning and surface sources. In the O₃ analysis, the influence of lightning is at least equal to that of surface emissions from June to November and is twice as large from July to September in the upper troposphere. The largest lightning impact on the upper tropospheric

O₃ budget reaches ≈30% in August. Below 5 km O₃ production is dominated by surface NO_x emissions, with the lightning contribution reaching 20% in August. Over the entire troposphere, lightning and surface NO_x emissions account for comparable O₃ production in July and August. Note that, in addition to ozone production caused by NO_x, transport and chemical processing also strongly affect tropospheric O₃, including stratospheric injection, horizontal advection, and *in situ* chemical production/destruction.

The horizontal extent of influence exerted by U.S. lightning is depicted in Fig. 3, showing the percentage changes in NO_x (*Right*) and O₃ (*Left*) concentrations at the 250-mbar level averaged for the month of July. At this level, NO_x concentrations over the U.S. and other large areas of the Northern Hemisphere are noticeably controlled by lightning. An area of 25–60% influence extends across Europe and into Asia. U.S. lightning also has a significant influence on O₃ concentrations for much of the Northern Hemisphere. There is 25–40% influence on the O₃ concentrations over the entire Gulf of Mexico and a large area of the Atlantic Ocean east of the U.S. Influence of at least 20% extends across the Atlantic Ocean into Europe.

The dominance of lightning in controlling summertime NO_x and O₃ in the middle and upper troposphere is unexpected, considering the fact that lightning accounts for only a small fraction of the total NO_x emission over the U.S. (7). The main difference between lightning and surface NO_x sources lies in the vertical distribution of the emissions. Lightning consists of both CG and IC flashes: a typical CG flash extends vertically up to 8 km, whereas a representative IC flash occurs at ≈6–10 km (16). Hence lightning directly deposits NO_x in the free troposphere. In contrast, NO_x emissions from anthropogenic activity, biomass burning, and soil are released within the atmospheric planetary boundary layer (PBL). Because the lifetime of NO_x varies considerably with altitude, being only a few hours near the PBL and up to a few days in the upper troposphere (17, 18), NO_x emitted directly in the mid and

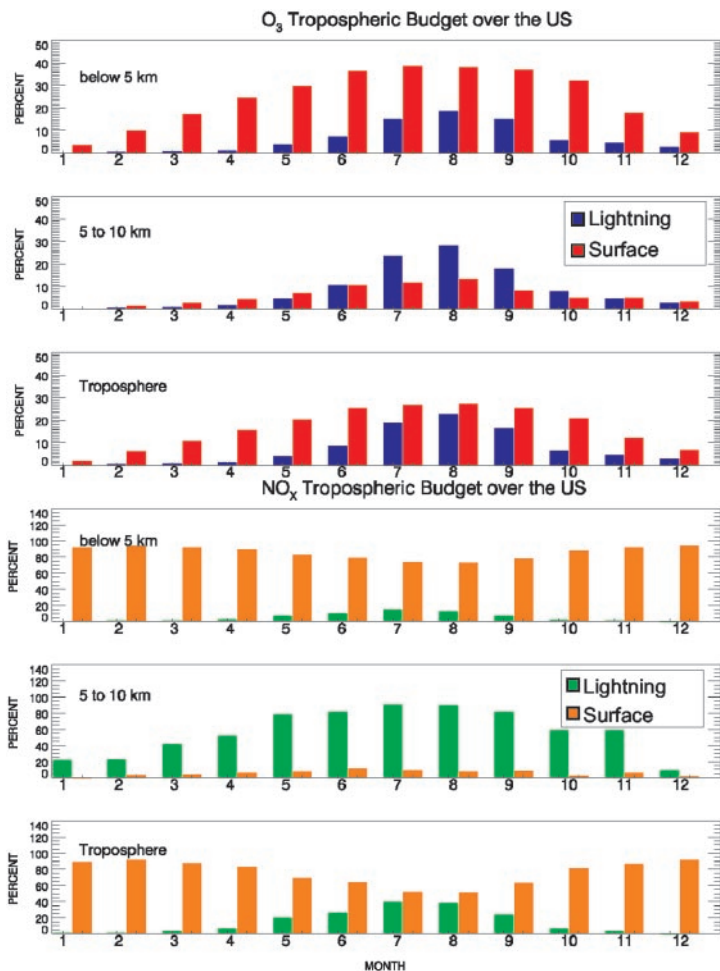


Fig. 2. Source influences on NO_x and O_3 budgets. The y axis represents the mass-weighted percentage change in concentrations summed vertically over the altitudes and horizontally over the contiguous U.S. domain ($30\text{--}50^\circ\text{ N}$ and $75\text{--}125^\circ\text{ W}$). In the NO_x panels, green columns represent lightning emission and orange columns represent the surface emissions (i.e., anthropogenic, biomass burning, and soil). In the O_3 panels, the lightning case is shown in blue and the surface emission case is in red. Months are shown on the x axis with 1 denoting January and 12 denoting December. The troposphere panels correspond to the altitudes between 0 and 10 km.

upper troposphere is more efficient in influencing NO_x and producing ozone in that region. In addition, emissions of lightning and surface NO_x sources are seasonally and geographically different. Lightning occurs primarily in the southeast portion of the U.S. in the summer, whereas anthropogenic NO_x is emitted in many urban areas throughout the country and the year. Surface NO_x emissions do not always coincide with the convective process. Our model calculations indicate that surface NO_x emissions indeed contribute to the NO_x and O_3 budgets in the upper troposphere (Fig. 3), consistent with the notion that convection associated with thunderstorms transports chemical compounds from the polluted PBL (19). However, such a process may experience a significant dilution with ambient air before reaching the mid or upper cloud region (20). Field and satellite observations show that near mature convective storms in the subtropical and mid-latitude regions lightning is primarily responsible for a significant NO_x enhancement in the upper troposphere (21–23).

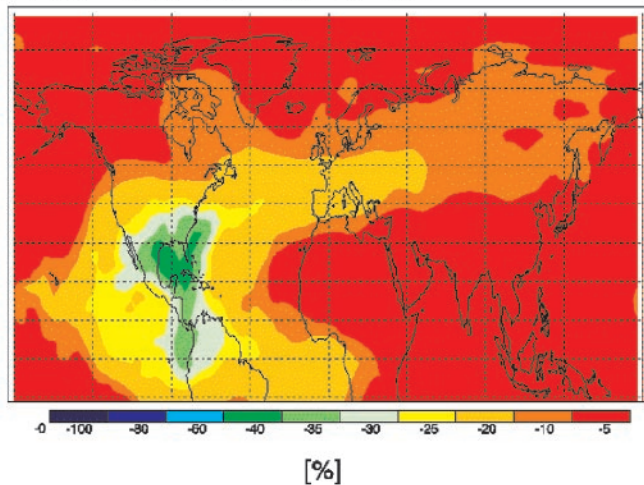
Several field campaigns in North America measured the concentrations of NO , NO_x , and O_3 (24). The field campaigns considered in the present study cover the following spatial and temporal ranges: Elchem, July 27–August 22, 1989, $30\text{--}35^\circ\text{ N}$ and $110\text{--}105^\circ\text{ W}$; Sterao, June 26–July 16, 1996, $39\text{--}42^\circ\text{ N}$ and $78\text{--}74^\circ\text{ W}$; Able3, July 6–August 15, 1990, $35\text{--}45^\circ\text{ N}$ and $80\text{--}70^\circ\text{ W}$; and Cite2, August 11–September 5, 1986, $35\text{--}45^\circ\text{ N}$ and $125\text{--}110^\circ\text{ W}$. The observed vertical mixing ratio profiles of NO , NO_x , and O_3 (averaged over the duration of the studies) are compared with the corresponding concentrations calculated by model runs for the same time period and location in Fig. 4.

Inclusion of lightning NO_x production in the model is essential to reproduce NO and NO_x enhancement in the upper troposphere. For the field campaign conducted along the heavily populated Northeast of the U.S. (Able3), the inclusion of surface NO_x sources is necessary to bring agreement between the modeled and measured NO_x concentrations in the lower atmosphere. The model simulations including both lightning and surface NO_x emissions produce O_3 concentrations that are consistent with the field observations, but removal of the surface emissions leads to an underpredicted O_3 in the lower troposphere. The good agreement between modeled and measured concentrations of NO_x , NO , and O_3 hence validates the treatment of the various NO_x emissions in our model simulations.

Conclusions

In this article we have assessed the impact of anthropogenic and natural NO_x sources over the contiguous U.S. on tropospheric NO_x and O_3 levels by using a global 3D chemical transport model. The results provide a quantitative assessment of the impacts of regional lightning over the U.S. on tropospheric chemistry, unambiguously establishing that summer lightning over the U.S. plays a surprising role in controlling NO_x and O_3 concentrations in the middle and upper troposphere. There have been intensive debates on NO_x emissions and control regulations among the scientific and industrial communities and at the political decision-making level (Intergovernmental Panel on Climate Change, Summary for Policymakers by Working Group 1, www.ipcc.ch). A reduction in fossil-fuel combustion leads to a

Changes in O₃ due to lightning at 250mb July



Changes in NO_x due to lightning at 250mb July

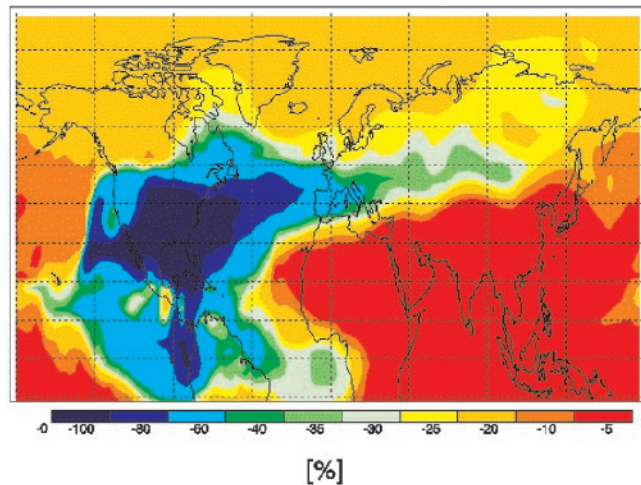


Fig. 3. Lightning impact on O₃ and NO_x at 250 mbar in July. Shown are percentage changes in NO_x (Right) and O₃ (Left) concentrations at the 250-mbar level averaged for the month of July.

lesser amount of tropospheric ozone precursors and can be essential to abate urban air pollution. In particular, the U.S. produces a significant fraction of the world anthropogenic NO_x emission ($\approx 33\%$). Our model calculations confirm that fossil-fuel combustion over the U.S. contributes overwhelmingly to the NO_x budget and exacerbates O₃ formation in the lower troposphere. On the other hand, the results reveal that summer lightning over the U.S. dominates the NO_x budget in the middle and upper troposphere and strongly affects regional and global

O₃ production. Hence a comprehensive assessment of the atmospheric NO_x emission sources and their impact is required to devise control strategies for regional and global air pollution. Furthermore, there is now increasing evidence that cloud electrification is inadvertently modified because of urbanization and air pollution (25), leading to significantly more lightning flashes. Such enhanced lightning could exert profound effects on the tropospheric oxidizing power, global biogeochemical cycles, and the climate.

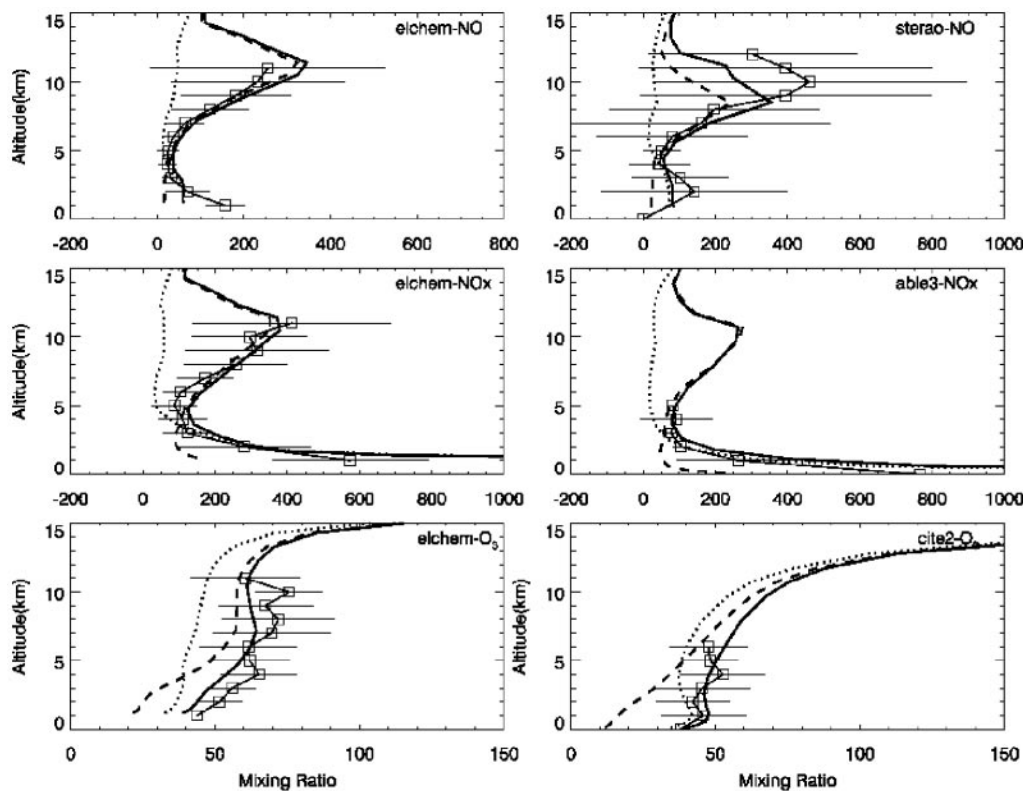


Fig. 4. Vertical mixing ratio profiles of NO, NO_x, and O₃ from model results and aircraft observations. The solid line with symbols shows the averages of concentrations observed from aircraft during the field studies. The horizontal lines represent the standard deviations of the observations. The solid line without symbols shows the concentrations calculated in model run *i*. Similarly, the dotted line represents model run *ii*, and the dashed line corresponds to model run *iii*.

We thank Prof. Robert A. Duce for helpful comments on this manuscript. This research was partially supported by the National Aeronautics and Space Administration New Investigator Program in

Earth Science and the Texas Air Research Center. The National Center for Atmospheric Research is supported by the National Science Foundation.

1. Crutzen, P. J. (1970) *Q. J. R. Meteorol. Soc.* **96**, 320–327.
2. Thompson, A. M. (1992) *Science* **256**, 1157–1165.
3. Seinfeld, J. H. & Pandis, S. N. (1998) *Atmospheric Chemistry and Physics: From Air Pollution to Global Change* (Wiley, New York).
4. Schlesinger, W. H. (1997) *Biogeochemistry* (Academic, New York).
5. Navarro-Gonzalez, R., Mckay, C. P. & Mvondo, D. N. (2001) *Nature* **412**, 61–63.
6. Ramanathan, V. & Dickinson, R. E. (1979) *J. Atmos. Sci.* **36**, 1084–1104.
7. Bond, W. D., Zhang, R., Tie, X., Brasseur, G., Huffines, G., Orville, R. E. & Boccippio, D. J. (2001) *J. Geophys. Res.* **106**, 27701–27710.
8. Lamarque, J. F., Brasseur, G. P., Hess, P. G. & Müller, J. F. (1996) *J. Geophys. Res.* **101**, 22995–22968.
9. Penner, J. E., Lister, D. H., Griggs, D. J., Dokken, D. J. & McFarland, M., eds. (1999) *Aviation and the Global Atmosphere: A Special Report of IPCC Working Groups I and III* (Cambridge Univ. Press, New York).
10. Olivier, J. G. J., Bouwman, A. F. J., Berdowski, J. M., Veldt, C., Bloos, J. P. J., Visschedijk, A. J. H., van der Maas, C. W. M. & Zandveld, P. Y. J. (1996) *Description of EDGAR Version 2.0: A Set of Global Emission Inventories of Greenhouse Gases and Ozone-Depleting Substances for All Anthropogenic and Most Natural Sources on a Per-County Basis and on a 1°1 Degree Grid, RIVM Report 771060 002/TNO-MEP Report R96/119* (National Institute for Public Health and the Environment, Bilthoven, The Netherlands).
11. Olivier, J. G. J., Bouwman, A. F., Van der Maas, C. W. M., Berdowski, J. J. M., Veldt, C., Bloos, J. P. J., Visschedijk, A. J. H., Zandveld, P. Y. J. & Haverlag, J. L. (1999) *Environ. Sci. Policy* **2**, 241–264.
12. Brasseur, G. P., Hauglustaine, D. A., Walters, S., Rasch, P. J., Müller, J. F., Granier, C. & Tie, X. (1998) *J. Geophys. Res.* **103**, 28265–28289.
13. Nesbitt, S. W., Zhang, R. & Orville, R. E. (2000) *Tellus B* **52**, 1206–1215.
14. Bond, D. W., Steiger, S., Zhang, R., Tie, X. & Orville, R. E. (2002) *Atmos. Environ.* **36**, 1509–1519.
15. Price, C., Penner, J. & Prather, M. (1997) *J. Geophys. Res.* **102**, 5929–5941.
16. Uman, M. A. (1987) *The Lightning Discharge* (Academic, San Diego).
17. Tie, X., Zhang, R., Brasseur, G., Emmons, L. & Lei, W. (2001) *J. Geophys. Res.* **106**, 3167–3178.
18. Tie, X., Zhang, R., Brasseur, G. & Lei, W. (2002) *J. Atmos. Chem.* **43**, 61–74.
19. Dickerson, R. R., Huffman, G. J., Luke, W. T., Nummermacker, L. J., Pickering, K. E., Leslie, A. C. D., Lindsey, C. G., Slinn, W. G. N., Kelley, T. J., Daum, P. H., *et al.* (1987) *Science* **235**, 460–464.
20. Telford, J. W. & Chai, S. K. (1993) *J. Appl. Meteorol.* **32**, 700–715.
21. Ridley, B. A., Dye, J. E., Ealega, J. G., Zheng, J., Grahek, F. E. & Rison, W. (1996) *J. Geophys. Res.* **101**, 20985–21005.
22. Huntrieser, H., Schlager, H., Feigl, C. & Holler H. (1999) *J. Geophys. Res.* **103**, 28247–28264.
23. Zhang, R., Sanger, N. T., Orville, R. E., Tie, X., Randel, W. & Williams, E. R. (2000) *Geophys. Res. Lett.* **27**, 685–688.
24. Bradshaw, J., Davis, D., Grodzinsky, G., Smyth, S., Newell, R., Sandholm, S. & Liu, S. (2000) *Rev. Geophys.* **38**, 61–116.
25. Orville, R. E., Huffines, G., Nielsen-Gammon, J., Zhang, R., Ely, B., Steiger, S., Phillips, S., Allen, S. & Read, W. (2001) *Geophys. Res. Lett.* **28**, 2597–2600.