

Dynamics of origination and extinction in the marine fossil record

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The discipline-wide effort to database the fossil record at the occurrence level has made it possible to estimate marine invertebrate extinction and origination rates with much greater accuracy. The new data show that two biotic mechanisms have hastened recoveries from mass extinctions and confined diversity to a relatively narrow range over the past 500 million years (Myr). First, a drop in diversity of any size correlates with low extinction rates immediately afterward, so much so that extinction would almost come to a halt if diversity dropped by 90%. Second, very high extinction rates are followed by equally high origination rates. The two relationships predict that the rebound from the current mass extinction will take at least 10 Myr, and perhaps 40 Myr if it rivals the Permo-Triassic catastrophe. Regardless, any large event will result in a dramatic ecological and taxonomic restructuring of the biosphere. The data also confirm that extinction and origination rates both declined through the Phanerozoic and that several extinctions in addition to the Permo-Triassic event were particularly severe. However, the trend may be driven by taxonomic biases and the rates vary in accord with a simple log normal distribution, so there is no sharp distinction between background and mass extinctions. Furthermore, the lack of any significant autocorrelation in the data is inconsistent with macroevolutionary theories of periodicity or self-organized criticality.

biodiversity | macroevolution | mass extinction

Decades of literature on large-scale taxonomic diversification and extinction patterns have hinged on compilations that record little more than first and last appearances of families or genera. Key examples include Sepkoski's compendia of marine families (1) and genera (2) and the Fossil Record 2 database of marine and continental families (3). Numerous patterns of widespread scientific and public interest have been identified on the basis of the older compilations, such as the identity of the five largest mass extinctions (4, 5), a gradual decline of extinction rates throughout the entire Phanerozoic (4), and possible cycles in extinction rates (6). A complete reevaluation of these hypotheses is now made possible by the maturation of the Paleobiology Database, a relational, web-based, and much more detailed resource created by and for the paleontological community (7).

Arguably, the most enduring and biologically important question these data can answer is whether global biodiversity is saturated (1, 8, 9). If so, then ecological interactions, such as competition and predation, must control rates of speciation and extinction (10–12). Speciation rates must be lower or extinction rates must be higher than they would be without these interactions. Diversity curves should increase logistically as they approach the saturation point (8) instead of exponentially (13). Increases in the diversity of major taxonomic groups should be balanced by decreases in the diversity of other groups (9). Most importantly, any recovery from a mass extinction, such as the current one, should eventually bring diversity back to the saturation point. Of course, the recovery will be rapid only in geological terms, the saturation point may change, and the extinction may fundamentally reorganize the global biota both taxonomically and ecologically, as seen in the wake of major perturbations, such as the end-Permian crisis (14).

Past predictions about recovery have been hampered by limited direct evidence for saturation in the fossil record.

Sepkoski (1, 8, 9) did argue in detail that turnover rates have constrained the global diversity of all marine animals over the entire Phanerozoic. Some studies of particular taxonomic groups over specific parts of the Phanerozoic also suggested density-dependent dynamics (15–19). However, both earlier (20) and later (13) workers argued that Phanerozoic diversity was not constrained. Even though this view is very inconsistent with such well documented patterns as rapid rebounds from mass extinctions (14, 21, 22), a basic logistic model assuming a single equilibrium point (8) failed to explain in a simple way why diversity appeared to increase exponentially in the Cretaceous and Cenozoic, toward the end of the Phanerozoic (1). Instead, a model with multiple equilibria was invoked (1, 9).

More recently, evidence has accumulated that the late Phanerozoic radiation is actually a combined artifact of increased sampling intensity and a related effect called the Pull of the Recent (7, 23–25). Thus, the question of diversity equilibrium has been reopened. The strong statistical patterns reported here show that diversity does not increase exponentially without constraints and therefore make it possible to predict the rebound from the current mass extinction in strict quantitative terms. First, however, a series of other major hypotheses concerning mass extinctions and diversity dynamics need to be addressed.

Pulsed Turnover Rates. Based on Sepkoski's classic family- and genus-level data (1, 2), it has been suggested that turnover comes in large pulses that coincide with interval boundaries (26, 27). If true, this result has the profound implication that even background turnover is largely forced either by perturbations, such as eruptions, sea level and climate changes, and bolide impacts (28), or by episodic ecological interactions, such as cascading extinctions (29). If turnover is not coupled with boundaries, an alternative hypothesis is that background extinction is effectively stochastic and results from the never ending process of competition over a fixed or slowly changing resource base, i.e., the Red Queen hypothesis (30).

The pulsed turnover hypothesis implies that per-million year (Myr) rates will correlate inversely with interval lengths because the assumption that turnover is continuous is violated (6). There is such a relationship for per-Myr extinction rates (Spearman rank-order correlation $\rho = -0.409$, $P < 0.005$). However, the same correlation does not exist in the unstandardized rates [$\rho = -0.021$, not significant (n.s.)] and is not significantly different from a distribution generated by bootstrapping (i.e., correlating raw rates with ratios of themselves to randomly drawn bin lengths). Thus, the relationship can be explained as resulting

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from random variation in the bin lengths because of random errors in the underlying time scale. The same pattern is seen with originations, i.e., first appearances of genera. Origination rates correlate negatively with bin length if they are standardized ($\rho = -0.408$, $P = 0.005$) but otherwise do not ($\rho = 0.155$, n.s.). These results, however, do not particularly endorse the Red Queen hypothesis because rates are still quite variable, as discussed below.

On balance, it is most parsimonious to infer that, although turnover may or may not be pulsed, interval boundaries do not coincide consistently with pulses, so continuous-time rates are more realistic than turnover proportions. These rates also have the advantage of removing the upper bound present in proportions, which can cause analytical artifacts. The continuous rates are not standardized for bin length in subsequent analyses because of the time scale's relative homogeneity and the suggestion that doing so would bias them. For the current dataset, this technical problem is most likely moot, because there is no trend through time in bin length ($\rho = 0.168$, n.s.), and the standard deviation of logged bin lengths is modest (0.348).

Phanerozoic Decline in Extinction Rates. Again based largely on Sepkoski's data, it has long been believed that there has been a decline through the Phanerozoic in both extinction rates (4) and origination rates (31). These observed declines are robust to the choice of rate metrics (26). Indeed, the new data clearly support a decline in both kinds of rates (extinction vs. time: $\rho = 0.547$, $P < 0.001$; origination vs. time: $\rho = 0.533$, $P < 0.001$). The patterns are influenced by extremely high values at the beginning of the time series that represent the Cambrian and earliest Ordovician (Fig. 1). However, removing these points does not greatly weaken the trends (extinction: $\rho = 0.446$, $P = 0.003$; origination: $\rho = 0.465$, $P = 0.002$). It is noteworthy that the correlations still appear even though the earlier studies (4, 31, 32) used finer time scales and therefore had greater statistical power. They also did not correct for sampling biases that would favor finding such a pattern.

To quantify the steepness of the declines, it is appropriate to perform a linear regression after log-transforming the turnover rates (26, 33), which is necessary because they are skewed and bounded by zero. For extinction and origination after the earliest Ordovician, the respective regression slopes are 0.201% and 0.158% per Myr, and the intercepts at 0 Ma are 0.218 and 0.282. Sepkoski's data imply much steeper slopes and predict much lower extinction rates for the Neogene (34), which is expected because the data are influenced by the Pull of the Recent.

The drop in rates could be explained in at least four ways. First, a trend might be created by sampling biases or an increase through the Phanerozoic in the average durations of sampling bins (35). The former problem has been fully resolved by sampling standardization of the data and by the use of rate equations that are robust to edge effects such as the Pull of the Recent [see [supporting information \(SI\) Text](#)]. The latter problem has been resolved by careful lumping of stages to produce bins of approximately equal length, as mentioned above.

Second, the trends could reflect a handoff between major groups that had coupled logistic dynamics (1). Groups that were dominant early on might have had higher intrinsic turnover rates but lower carrying capacities, leading eventually to their replacement by slowly radiating competitors. Here, these groups would be the three "evolutionary faunas" (36). Evidence is lacking for dynamically distinct but interacting groups of taxa based on actual turnover rates in Sepkoski's own genus-level compendium (37), so the evolutionary fauna hypothesis is not relevant here and is most likely useful only to summarize coincidences between diversity trends (36) and onshore-offshore patterns (38).

Third, individual orders with high turnover rates might have randomly gone extinct by the mid-Paleozoic simply because of

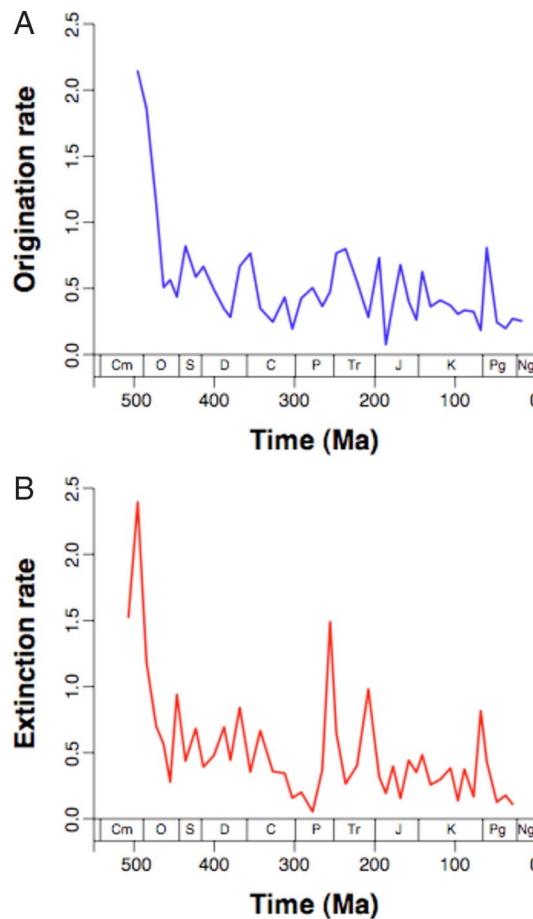


Fig. 1. Per-interval instantaneous origination rates (A) and extinction rates (B) of marine invertebrate genera over the Phanerozoic. Data are binned into 48 intervals averaging 11.0 Myr in duration.

their volatile dynamics, leaving only the groups with low turnover rates (32). Neither this hypothesis nor the preceding one can be tested directly by using the new data without separating the major taxonomic groups, but at least the volatility model is supported by analyses of Sepkoski's data (32).

Finally, higher taxa such as families may tend to accumulate more species through time, reducing their chance of extinction (39). A similar argument holds for origination. Raup (40) provided estimated species totals for the Cenozoic and for each Paleozoic and Mesozoic period that can be compared with genus totals extracted from Sepkoski's compendium (2). Although the ratios are probably too low, because the species counts are two decades older, there is a striking increase from 2.6, 2.7, and 2.1 in the Cambrian, Ordovician, and Silurian, to 5.4 and 7.8 in the Cretaceous and Cenozoic, with all other periods falling in between. More detailed analyses are called for, but the difference is so large that it could easily explain the trend.

Big Five Mass Extinctions. The apparent existence of five major peaks in extinction rates is another key finding in the literature on Phanerozoic marine diversity (4, 5). The Big Five theory is so widespread that it has given rise to the popular term "sixth extinction" in relation to the current crisis (41, 42). These peaks were first recognized because they rose beyond the parametric 95% confidence interval around a linear regression that described the Phanerozoic decline in extinction rates (4). The Big Five also appear to be present in a separate compilation of marine family-level data (13), and at least the Permo-Triassic

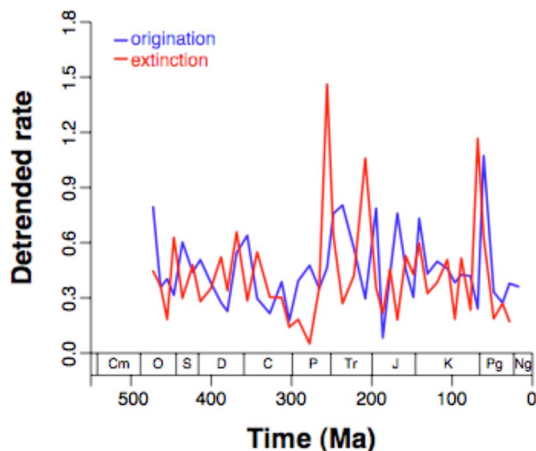


Fig. 2. Origination rates (blue lines) and extinction rates (red lines) after being detrended by using exponential functions.

and Triassic–Jurassic events register as well in data on continental organisms (13).

However, there were numerous problems with the original analysis, including failure to log the data (33) and the use of per-Myr but not per-taxon rates. Indeed, the original data (4) did not particularly support the existence of two mass extinctions now widely discussed, the late Devonian and end-Triassic (43). Although the five greatest proportional drops in diversity in Sepkoski's later genus-level data (2) do match the conventional Big Five extinctions, the late Devonian and end-Triassic drops apparently are driven more by origination deficits than by high extinction rates (ref. 44; however, in the new data, the latter decline is a legitimate mass extinction). Furthermore, the end-Ordovician peak in Sepkoski's family-level data has been called into question because its appearance depends on the rate metric being used (26).

Finally, the very idea that the Big Five represent anything but the upper end of a continuous range of variation has been questioned almost from the start (5, 45). Although there are other good ways to test for outliers in rate distributions (45), it is sufficient to compare the rates with what is expected of the same number of observations drawn from a best-fit normal distribution. The rates first need to be detrended by taking residuals of the regression line that was earlier fit to the logged data. The detrended distributions (Fig. 2) are in fact indistinguishable from the log normal based either on a Kolmogorov–Smirnov test (extinction: $D = 0.0888$, n.s.; origination: $D = 0.0883$, n.s.) or a Shapiro–Wilk test (extinction: $W = 0.9657$, n.s.; origination: $W = 0.9548$, n.s.). Additionally, although extinction rates are much more variable than origination rates in Sepkoski's datasets (26), here the standard deviations of the detrended rates are very similar (extinction: 0.613; origination: 0.466) and the distributions are statistically indistinguishable ($D = 0.2215$, n.s.). Thus, not only are there no distinct classes of major origination or mass extinction events, but there is minimal evidence that extinction rates are more volatile than origination rates.

These results do not challenge the idea that true mass extinctions impose different rules of survivorship relating to factors such as geographic range size (46), because the rules could well vary with extinction intensity even if intensity *per se* is largely random. They also do not challenge the idea that we are in the midst of a mass extinction on par with all but a handful of such events over the last half billion years (47).

Furthermore, the data confirm that three of the Big Five extinctions (the Permo-Triassic, end-Triassic, and Cretaceous-Tertiary) are the three largest of the Phanerozoic, if only after

detrending the rates (Fig. 2). These events stand out very clearly if the data are not log-transformed. Additionally, although the Paleozoic data show very little variability, one of the two largest events within the Paleozoic is indeed the end-Ordovician extinction. The other comes at the Devonian–Carboniferous boundary, immediately after the late Devonian event typically ranked in the Big Five. As mentioned, the late Devonian episode's importance already has been called into question (26, 43, 44). In summary, it is a matter of taste whether to speak of the Big Five, the Big Three, or just the Big One, but one way or another major mass extinctions are truly very rare.

Periodicity and Self-organized Criticality. The hypothesis that mass extinctions show an ≈ 26 -Myr cycle of periodicity (6) has arguably been the most debated hypothesis by paleontologists over the last quarter century. A host of potential mechanisms for periodicity have been offered, such as impacts driven by astrophysical factors (6) or the lag time needed either to accumulate species vulnerable to mass extinction (48) or to build networks of ecological interactions that can collapse if even slightly perturbed (29).

Qualitatively, it is hard to discern some of the extinction rate peaks that should fall in the periodic pattern (13). Quantitatively, extinction rates in the Fossil Record 2 family data (3) and Sepkoski's family and genus data (1, 2) are not correlated with themselves at any time lag (49), which is a necessary condition for periodicity to hold. That said, analyses of origination rates in all three datasets (49, 50) suggest short-term autocorrelation. However, the current dataset shows no autocorrelation in either kind of rate (Fig. S1), and a standard spectral analysis (Fig. S2) also suggests purely random variation through the time series (i.e., white noise).

The idea that mass extinctions may result from small perturbations of complex ecosystems (29) has been expanded to argue that food webs naturally evolve to a state of self-organized criticality that creates the extinction cascades, even in the absence of perturbations (51, 52). Criticality models make direct predictions about patterns of autocorrelation in time series (52, 53). However, results suggesting criticality in several major paleontological datasets (52) were quickly rebutted as statistical artifacts that resulted from interpolating turnover rates to create numerous evenly spaced intervals (54, 55). With the new data, the slope of the spectral density/frequency relationship is far from the value predicted by criticality; it is effectively zero (Fig. S2B). Thus, self-organized criticality can be dismissed as a plausible explanation of extinction rates.

Finally, it also has been suggested that diversity itself is cyclical with a period of ≈ 62 Myr, regardless of turnover rates (56). This pattern was again compromised by the use of Sepkoski's unstandardized data, and it had serious plausibility problems: Although extinction rates do not show long-range periodicity, most of the predicted peaks only appeared to be peaks because they were followed by the nominal Big Five extinctions (e.g., 4, 49). No distinct peaks were seen within the Cretaceous and Cenozoic (56) (Fig. S3), and there is also no evidence in the current dataset (25) for the predicted early Cambrian peak, or for a valley separating the supposed late Ordovician and Devonian peaks.

The current diversity curve does show a weak cycle after removing its U-shaped trend by taking residuals of a quadratic function fitted to the logged data (Fig. S3). However, the cycle has a period somewhat longer than 62 Myr (Fig. S1C), and the pattern is driven by clearly coincidental matches between two peaks and two valleys (Fig. S3). The spectral density data (Fig. S2C) also suggest periodicity. However, the steep falloff at medium frequencies is driven by these weak matches and consistent with either a power law or exponential function, so it is not evidence of self-organized criticality (55).

Previous Studies. Put together, these results combine to make a plausible and straightforward scenario. First, a major drop in diversity for any reason will be followed by a significant recovery due to the near-absence of extinction. Extinction rates may continue to be low for a second 11.0-Myr interval after the immediate recovery. Second, if the drop was due to a large extinction pulse and not merely an unfavorable balance of origination and extinction, then speciation rates will increase dramatically.

Earlier researchers have argued for entirely different models. For example, autocorrelation in Sepkoski's origination rates (49); apparently greater variation in extinction than origination (50); and, most importantly, a lagged correlation between extinction and subsequent origination in Sepkoski's family- and genus-level data (21) have been used to construct a theory that diversity is slow to rebound from extinction because time is required to reconstruct ecological niches that might be filled.

There are numerous problems with Sepkoski's data that call these results into question: the lack of any sampling standardization, the unavoidable backwards smearing of extinction rates and forward smearing of origination events that results from using simple range data (22), and the oversplitting of the time scale into 106 intervals (as opposed to 48 in this study). Not surprisingly, the first two hypothesized patterns do not exist in the current dataset: Origination rates lack any significant autocorrelation and are not much less variable than extinction rates (Figs. S1 and S2).

An extinction-origination correlation is indeed present, but the pattern is different from expected under the hypothesis that niches need to be reconstructed. First, very high origination rates come immediately after what are clearly rapid mass extinctions (Fig. 3B), not a full temporal bin later (21). Second, only the largest extinctions seem to boost immediately subsequent origination. If Kirchner and Weil (21, 49, 50) are right that niches need to be reconstructed, then recoveries should be slower and not faster after a large extinction, because the relevant ecological interactions are more highly disrupted. Thus, if niches are relevant, then speciation in the wake of mass extinctions is most likely fostered by the lack of competition for existing niches instead of being delayed by the elimination of old niches that need to be reconstructed.

Meanwhile, some support for a diversity/extinction relationship has been found in Sepkoski's datasets (1, 57). However, these results did not involve lagging, and Sepkoski's genus-level data provide mixed support for density dependence in both kinds of rates instead of just extinction rates (57). The consistency of the diversity-extinction relationship through the Phanerozoic (Fig. 3A) also is at odds with earlier analyses suggesting fundamental changes across the Permo-Triassic in ecology (14, 61) and especially diversity dynamics (57). The latter study found correlations between changes in diversity and changes in rates. The motivation for differencing the rates was to avoid biases and autocorrelation (57), but the current dataset uses rates with low bias that lack autocorrelation (Figs. S1 and S2), so differencing is moot. Again, biases related to sampling, counting, and rate equations are likely to be a problem in all of the earlier studies.

Recovery Predictions. The two major correlational relationships (Fig. 3) are well constrained over a realistic range of diversity and turnover levels, making it reasonable to offer specific predictions about the recovery from the mass extinction that is clearly underway (47, 62). First, however, the relationships need to be modeled as accurately as possible, which requires transforming the data appropriately and then fitting linear regression functions.

The detrended diversity curve (Fig. S3) falls in such a narrow range that it is normally distributed on either a linear, log, or square root scale according to a Shapiro-Wilk test. However, logging diversity data is intuitive because diversification is a

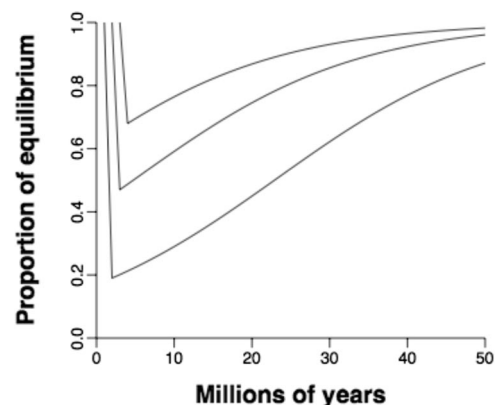


Fig. 4. Predicted recoveries from mass extinctions modeled on the global Permo-Triassic and Cretaceous-Tertiary events (81% and 53% losses, respectively) and the early Pleistocene extinction in the Caribbean (32% loss) (67).

multiplicative process. For the detrended extinction rates, the raw data are far too skewed to be normal ($P < 0.001$), and taking square roots does not really help ($P = 0.030$). However, logging does render them normal, and the parametric correlation r between log diversity and log future extinction (0.462) is much the same as the rank-order correlation ρ (0.439). Like extinction rates, origination rates are too skewed to be normal ($P = 0.011$). Technically, logging does normalize them ($P = 0.090$), but the square root transform does a slightly better job ($P = 0.455$).

For the data starting after the earliest Ordovician, the median extinction rate is 0.380, and the diversity:extinction regression line is so strong that it implies a near-zero rate when diversity is 1 (i.e., the intrinsic rate). For example, at diversity levels 50, 90, and 99% below the median, the predicted extinction rates are 0.133, 0.014, and 0.0005, respectively. Thus, there would be hardly any extinction if not for major environmental perturbations and the ecological interactions that generate density dependence. Of course, that does not mean the rates are entirely predictable; it simply means that if abundant data were to extend all the way down to a diversity level of 1, we would find that density dependence explained a large majority of the variance. Indeed, the residual variance of the actual regression (Fig. 3A) is substantial and, therefore, likely to be real.

The second step is to model origination as a function of past extinction. However, origination rates cannot be predicted solely from the contemporary extinction rates produced in the first step without producing pathological results, because the initially low extinction rates would imply low, not high, origination rates. Instead, the extinction rates put into the equation need to reflect the entire loss of diversity relative to the starting point. The solution is to add the log ratio of preextinction diversity to current diversity to the predicted rate.

Put together, the two functions paint a grim picture (Fig. 4). A mass extinction on the scale of the Permo-Triassic event would probably leave diversity still 20% below its equilibrium level after ≈ 40 Myr, which is nearly as long as a typical geological period. Indeed, a comparable recovery from even the weakest modeled extinction is expected to take ≈ 10 Myr, and a 90% recovery would take ≈ 20 Myr. The worst-case scenario is not unthinkable for marine invertebrates: Any increase in global atmospheric CO_2 by >500 ppm would cause coral reef ecosystems to collapse (63), and, putting everything else aside, biotic homogenization through species introductions could cause up to a 58% mass extinction (64).

Origination rates do vary significantly from one 11.0-Myr bin to the next (Figs. 1A and 2A and Fig. S1A), so stochastically achieved high rates in a few bins could push diversity to

1. Sepkoski JJ, Jr (1984) A kinetic model of Phanerozoic taxonomic diversity. III. Post-Paleozoic families and mass extinctions. *Paleobiology* 10:246–267.
2. Sepkoski JJ, Jr (1996) in *Global Events and Event Stratigraphy*, ed Walliser OH (Springer, Berlin), pp 35–52.
3. Benton MJ (1993) *The Fossil Record 2* (Chapman & Hall, London).
4. Raup DM, Sepkoski JJ, Jr (1982) Mass extinctions in the marine fossil record. *Science* 215:1501–1503.
5. Raup DM (1986) Biological extinction in earth history. *Science* 231:1528–1533.
6. Raup DM, Sepkoski JJ, Jr (1984) Periodicity of extinctions in the geologic past. *Proc Natl Acad Sci USA* 81:801–805.
7. Alroy J, et al. (2001) Effects of sampling standardization on estimates of Phanerozoic marine diversification. *Proc Natl Acad Sci USA* 98:6261–6266.
8. Sepkoski JJ, Jr (1978) A kinetic model of Phanerozoic taxonomic diversity. I. Analysis of marine orders. *Paleobiology* 4:223–251.
9. Sepkoski JJ, Jr (1979) A kinetic model of Phanerozoic taxonomic diversity. II. Early Phanerozoic families and multiple equilibria. *Paleobiology* 5:222–251.
10. MacArthur RH (1969) Patterns of communities in the tropics. *Biol J Linn Soc* 1:19–30.
11. Rosenzweig ML (1975) in *Ecology and Evolution of Communities*, eds Cody ML, Diamond JM (Belknap, Cambridge, MA), pp 121–140.
12. Walker TD, Valentine JW (1984) Equilibrium models of evolutionary species diversity and the number of empty niches. *Am Nat* 124:887–899.
13. Benton MJ (1995) Diversification and extinction in the history of life. *Science* 268:52–58.
14. Erwin DH (2001) Lessons from the past: Biotic recoveries from mass extinctions. *Proc Natl Acad Sci USA* 98:5399–5403.
15. Mark GA, Flessa KW (1977) A test for evolutionary equilibria: Phanerozoic brachiopods and Late Cenozoic New World mammals. *Paleobiology* 3:17–22.
16. Wagner PJ (1995) Diversity patterns among early gastropods: Contrasting taxonomic and phylogenetic descriptions. *Paleobiology* 21:410–439.
17. Alroy J (1996) Constant extinction, constrained diversification, and uncoordinated stasis in North American mammals. *Palaeogeogr Palaeoclimatol Palaeoecol* 127:285–311.
18. Alroy J (1998) in *Biodiversity Dynamics: Turnover of Populations, Taxa, and Communities*, eds McKinney ML, Drake JA (Columbia Univ Press, New York), pp 232–287.
19. Connolly SR, Miller AI (2002) Global Ordovician faunal transitions in the marine benthos: Ultimate causes. *Paleobiology* 28:26–40.
20. Flessa KW, Levinton JS (1975) Phanerozoic diversity patterns: Tests for randomness. *J Geol* 83:239–248.
21. Kirchner JW, Weil A (2000) Delayed biological recovery from extinctions throughout the fossil record. *Nature* 404:177a–180a.
22. Foote M (2003) Origination and extinction through the Phanerozoic: A new approach. *J Geol* 111:125–148.
23. Foote M (2000) Origination and extinction components of taxonomic diversity: General problems. *Paleobiology* 26 (suppl):74a–102a.
24. Peters SE, Foote M (2001) Biodiversity in the Phanerozoic: A reinterpretation. *Paleobiology* 27:583–601.
25. Alroy J, et al. (2008) Phanerozoic trends in the global diversity of marine invertebrates. *Science* 321:97–100.
26. Foote M (1994) Temporal variation in extinction risk and temporal scaling of extinction metrics. *Paleobiology* 20:424–444.
27. Foote M (2005) Pulsed origination and extinction in the marine realm. *Paleobiology* 31:6–20.
28. Raup DM (1992) Large-body impact and extinction in the Phanerozoic. *Paleobiology* 18:80–88.
29. Plotnick RE, McKinney ML (1993) Ecosystem organization and extinction dynamics. *Palaios* 8:202–212.
30. Van Valen L (1973) A new evolutionary law. *Evol Theory* 1:1–30.
31. Gilinsky NL, Bambach RK (1987) Asymmetrical patterns of origination and extinction in higher taxa. *Paleobiology* 13:427–445.
32. Gilinsky NL (1994) Volatility and the Phanerozoic decline of background extinction rates. *Paleobiology* 20:445–458.
33. Quinn JF (1983) Mass extinctions in the fossil record. *Science* 219:1239–1240.
34. Peters SE (2006) Genus extinction, origination, and the durations of sedimentary hiatuses. *Paleobiology* 32:387–407.
35. Pease CM (1992) On the declining extinction and origination rates of fossil taxa. *Paleobiology* 18:89–92.
36. Sepkoski JJ, Jr (1981) A factor analytic description of the Phanerozoic marine fossil record. *Paleobiology* 7:36–53.
37. Alroy J (2004) Are Sepkoski's evolutionary faunas dynamically coherent? *Evol Ecol Res* 6:1–32.
38. Sepkoski JJ, Jr, Miller AI (1985) in *Phanerozoic Diversity Patterns: Profiles in Macroevolution*, ed Valentine JW (Princeton Univ Press, Princeton), pp 153–190.
39. Flessa KW, Jablonski D (1985) Declining Phanerozoic background extinction rates: Effect of taxonomic structure? *Nature* 313:216–218.
40. Raup DM (1976) Species diversity in the Phanerozoic: A tabulation. *Paleobiology* 2:279–288.
41. Leakey RE, Lewin R (1995) *The Sixth Extinction: Patterns of Life and the Future of Humankind* (Doubleday, New York).
42. Glavin T (2007) *The Sixth Extinction: Journeys Among the Lost and Left Behind* (Thomas Dunne Books, New York).
43. Raup DM, Sepkoski JJ, Jr (1983) Mass extinctions in the fossil record. *Science* 219:1240–1241.
44. Bambach RK, Knoll AH, Wang SC (2004) Origination, extinction, and mass depletions of marine diversity. *Paleobiology* 30:522–542.
45. Wang SC (2003) On the continuity of background and mass extinction. *Paleobiology* 29:455–467.
46. Jablonski D (1986) Background and mass extinctions: The alternation of macroevolutionary regimes. *Science* 231:129–133.
47. Myers N, Knoll AH (2001) The biotic crisis and the future of evolution. *Proc Natl Acad Sci USA* 98:5389–5392.
48. Stanley S (1990) Delayed recovery and the timing of mass extinctions. *Paleobiology* 16:401–414.
49. Kirchner JW, Weil A (2000) Correlations in fossil extinction and origination rates through geological time. *Proc R Soc London Ser B* 267:1301–1309.
50. Kirchner JW (2002) Evolutionary speed limits inferred from the fossil record. *Nature* 415:65–68.
51. Snekken K, Bak P, Flyvbjerg H, Jensen MH (1995) Evolution as a self-organized critical phenomenon. *Proc Natl Acad Sci USA* 92:5209–5213.
52. Solé R, Manrubia SC, Benton M, Bak P (1997) Self-similarity of extinction statistics in the fossil record. *Nature* 388:764–767.
53. Solé R, Bascompte J, Manrubia SC (1996) Extinction: Bad genes or weak chaos? *Proc R Soc London Ser B* 263:1407–1413.
54. Kirchner JW, Weil A (1998) No fractals in fossil extinction statistics. *Nature* 395:337–338.
55. Newman MEJ, Eble GJ (1999) Power spectra of extinction in the fossil record. *Proc R Soc London Ser B* 266:1267–1270.
56. Rohde RA, Muller RA (2005) Cycles in fossil diversity. *Nature* 434:208–210.
57. Foote M (2000) Origination and extinction components of taxonomic diversity: Paleozoic and post-Paleozoic dynamics. *Paleobiology* 26:578b–605b.
58. Webb SD (1969) Extinction-origination equilibria in late Cenozoic land mammals of North America. *Evolution* 23:688–702.
59. Freckleton RP, Watkinson AR, Green RE, Sutherland WJ (2006) Census error and the detection of density dependence. *J Anim Ecol* 75:837–851.
60. Solé R, Bascompte J (1996) Are critical phenomena relevant to large-scale evolution? *Proc R Soc London Ser B* 263:161–168.
61. Wagner PJ, Kosnik MA, Lidgard S (2006) Abundance distributions imply elevated complexity of post-Paleozoic marine ecosystems. *Science* 314:1289–1292.
62. Wake DB (2008) Are we in the midst of the sixth mass extinction? A view from the world of amphibians. *Proc Natl Acad Sci USA* 105(Suppl):11466–11473.
63. Hoegh-Guldberg O, et al. (2007) Coral reefs under rapid climate change and ocean acidification. *Science* 318:1737–1742.
64. McKinney ML (1998) On predicting biotic homogenization: Species-area patterns in marine biota. *Global Ecol Biogeogr Lett* 7:297–301.
65. Jablonski D (2000) Micro- and macroevolution: Scale and hierarchy in evolutionary biology and paleobiology. *Paleobiology* 26 (suppl):15–52.
66. Alroy J (2000) New methods for quantifying macroevolutionary patterns and processes. *Paleobiology* 26:707–733.
67. Todd JA, et al. (2002) The ecology of extinction: Molluscan feeding and faunal turnover in the Caribbean Neogene. *Proc R Soc London Ser B* 269:571–577.
68. Payne JL (2005) Evolutionary dynamics of gastropod size across the end-Permian extinction and through the Triassic recovery interval. *Paleobiology* 31:269–290.
69. Raup DM (1985) Mathematical models of cladogenesis. *Paleobiology* 11:42–52.
70. Foote M (1999) Morphological diversity in the evolutionary radiation of Paleozoic and post-Paleozoic crinoids. *Paleobiology* 25 (suppl):1–115.
71. Alroy J (2008) in *Ecology and Speciation*, eds Butlin R, Bridle J, Schluter D (Cambridge Univ Press, Cambridge), in press.