Several archaeological studies in the Central Andes have pointed at the temporal coincidence of climatic fluctuations (both long- and short-term) and episodes of cultural transition and changes of socioeconomic structures throughout the pre-Cambrian period. Although most scholars explain the connection between environmental and cultural changes by the impact of climatic alterations on the capacities of the ecosystems inhabited by pre-Cambrian cultures, direct evidence for assumed demographic consequences is missing so far. In this study, we address directly the impact of climatic changes on the spatial population dynamics of the Central Andes. We use a large dataset of pre-Cambrian mitochondrial DNA sequences from the northern Rio Grande de Nasca drainage (RGND) in southern Peru, dating from ~840 BC to 1450 AD. Alternative demographic scenarios are tested using Bayesian serial coalescent simulations in an approximate Bayesian computational framework. Our results indicate migrations from the lower coastal valleys of southern Peru into the Andean highlands coincident with increasing climate variability at the end of the Nasca culture at ~640 AD. We also find support for a back-migration from the highlands to the coast coincident with droughts in the southeastern Andean highlands and improvement of climatic conditions on the coast after the decline of the Wari and Tiwanaku empires (~1200 AD), leading to a genetic homogenization in the RGND and probably southern Peru as a whole.

The role of transient or long-term climatic changes and resulting environmental impacts on human dispersals and demographic development is generally accepted (1, 2). Nevertheless, the exact mechanisms underlying such changes and the extent of their influence on local population histories of prehistoric humans remain poorly understood. Pre-Cambrian Andean cultures can offer exceptional insights into such human–environment interactions. In fact, western South America underwent several climate changes following the Pleistocene–Holocene transition since the initial peopling ~14,000 y ago (3, 4). The frequency of climatic changes seems to have been accelerated with the establishment of modern atmospheric conditions and increasing variability of the El Niño Southern Oscillation around the end of the Mid-Holocene (4, 5), resulting in multiple transient and protracted environmental impacts on the Central Andean region (6, 7). Several studies have pointed at the temporal coincidence of occasional environmental and catastrophic events (7, 8) and long-term climatic alterations (5, 9) with episodes of cultural transition and changes of socioeconomic structure throughout the pre-Cambrian period, suggesting that climate has played a major role in the formation of the cultural and demographic complexity of the Central Andean area (5, 7, 8, 10). Adaptive responses of the Andean populations to environmental impacts (e.g., droughts) such as the intensification of irrigation and extensive terracing have been directly inferred from the archaeological record (7, 8, 10–12). However, demographic changes including migration in response to climatic alterations have so far only been inferred indirectly (7, 9, 12, 13).

The heartland of most pre-Cambrian coastal cultures, especially in southern Peru, was located in the river oases of the lower valleys of the western Andean foothills. These settlement areas can be described as low-resilience, fragile ecosystems (5). Situated on the boundary between the hyperarid coastal desert and the high-altitude Andes, slight hydrological oscillations can change the availability of water, thereby triggering the fertility of the landscape. Interestingly, these ecosystems were inhabited very early in Andean prehistory (14, 15). Today there is nearly no direct precipitation in the western Andean foothills of southern Peru. The main sources of water are the rivers fed by seasonal rainfall in the headwaters of the high-altitude Andes. Recent investigations in the area around the modern city of Palpa, situated in the northern Rio Grande de Nasca drainage (RGND) area in southern coastal Peru (Fig. 1), have shown that the climatic conditions in this settlement area oscillated between humid and dry conditions in the Late Holocene (5). The RGND region received a reliable level of precipitation from the Early Horizon (EH; 840–260 BC) throughout the Early Intermediate Period (EIP; 260 BC–640 AD), culturally characterized by the Andean Horizon (EH; 840–260 BC) throughout the Early Intermediate Period (EIP; 260 BC–640 AD), culturally characterized by the

Significance

It has long been assumed that climate played a major role in the population history of the Central Andes. Although adaptations of the Andean populations to climatic changes such as the intensification of agriculture have been inferred from the archaeological record, evidence for demographic adaptations such as migration is missing so far. In this paper, ancient DNA data from populations that lived in southern Peru between 840 BC and 1450 AD provide evidence for two large-scale migrations in the Central Andes coincident with episodes of drought and increased climatic variability. These migrations led to a successive genetic homogenization of southern Peruvian populations generally attributed to intrusions by the late pre-Cambrian highland empires such as the Wari, Tiwanaku, or Inca.


This article is a PNAS Direct Submission.

Data deposition: The sequences reported in this paper have been deposited in the GenBank database (accession nos. KJ489100–KJ489306).

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successive climate thriving of Paracas and Nasca cultures (5, 15). According to local climate records, the variability of precipitation increased in the Late EIP (~600 AD), resulting in frequent droughts and occasional flash floods (16, 17). During the Middle Horizon (MH; 640–1180 AD), climatic conditions continued to become more unfavorable, peaking in a hyperarid phase starting around 750 AD (5). These local observations match those of a high-resolution marine climate record from the Peruvian shelf west of Lima (18), which indicates episodes of extreme drought between 800 and 1250 AD, following a peak in El Niño activity with heavy rainfalls. The climatic changes are coincident with a much lower settlement density in the lower valleys of the RGND (12, 19–22). It has been hypothesized that the inhabitants of the lower RGND valley successively retreated to the upper valleys of the settlement area because of the increasing aridity (5, 12, 13, 18). The increase of grass pollen from a peatland record 50 km east of Palpa shows that this humid period was precluded by a successive improvement of climatic conditions starting around 1150 AD in the northern RGND (17). The comparison of the Palpa climate record with ice-accumulation records from the Quelccaya ice cap in the southeastern Peruvian highlands (26, 27) and lake-level records of Lake Titicaca (10, 28) reveals that the oscillations between dry and humid conditions at the coast alternated with humid and dry spells in the southeastern Peruvian highlands (5). Contemporaneous with the coastal droughts during the Early MH, the area around Lake Titicaca saw the rise of the Tiwanaku culture and intensification of agriculture coinciding with more humid conditions (7, 10, 13). In the south-central highlands around Ayacucho, the rise and expansion of the Wari highland empire also began at the start of the MH (29–31). A period of severe aridity and cooler temperatures beginning in the early to mid-12th century (6, 32) has been identified as the proximate cause of the disintegration of the Tiwanaku state (33), contemporaneous with the period of increased humidity in the foothills of the RGND (5) and the demise of the Wari state (29).

Here we address directly the impact of climatic changes on the spatial population dynamics of the Central Andes by recording changes in the paleogenetic record with marked climate fluctuations and cultural transitions in the past. We performed population genetic analysis on a large dataset of published (34, 35) and newly obtained pre-Columbian mitochondrial DNA sequences from sites in the northern RGND area, spanning ~2,000 y. A series of alternative demographic scenarios was tested using serial coalescent analysis in an approximate Bayesian computational (ABC) framework.

Results and Discussion

Ancient DNA Analysis. We obtained mitochondrial (mt) hyper-variable region I (HVR I) sequences from 207 individuals of the northern RGND area deriving from coastal (coast and foothills) and highland (upper valleys) archaeological sites dating from the EH to the LIP. Of these, 157 HVR I sequences (mt haplotypes) had been published in previous studies (34, 35). We also reproducibly determined the mt haplotypes of 50 individuals from the coastal sites Monte Grande (EIP: n = 11; LIP: n = 11), Chillo (LIP: n = 11), and Los Molinos (MH: n = 5), and Tranca (MH: n = 12) in the highlands, to complement the geographical coverage in our research area (Fig. 1). The 207 sequences have been grouped into six populations depending on their chronological and ecogeographical origin (Table 1 and Dataset S1). The authenticity of the ancient DNA typing results has been ensured following the strict precautions common

Table 1. Summary genetic diversity statistics and mitochondrial haplogroup frequencies for the six analyzed ancient populations

<table>
<thead>
<tr>
<th>Population*</th>
<th>n</th>
<th>H †</th>
<th>Hd</th>
<th>π</th>
<th>Td</th>
<th>A</th>
<th>B</th>
<th>C</th>
<th>D</th>
</tr>
</thead>
<tbody>
<tr>
<td>C_EH</td>
<td>31</td>
<td>17</td>
<td>0.9527</td>
<td>0.0088</td>
<td>−1.208</td>
<td>0.06</td>
<td>0.00</td>
<td>0.20</td>
<td>0.74</td>
</tr>
<tr>
<td>C_EIP</td>
<td>66</td>
<td>42</td>
<td>0.9000</td>
<td>0.0117</td>
<td>−1.196</td>
<td>0.03</td>
<td>0.14</td>
<td>0.33</td>
<td>0.50</td>
</tr>
<tr>
<td>C_MH</td>
<td>11</td>
<td>9</td>
<td>0.9455</td>
<td>0.0127</td>
<td>−0.663</td>
<td>0.00</td>
<td>0.18</td>
<td>0.36</td>
<td>0.46</td>
</tr>
<tr>
<td>C_LIP</td>
<td>22</td>
<td>16</td>
<td>0.9654</td>
<td>0.0137</td>
<td>−0.589</td>
<td>0.05</td>
<td>0.59</td>
<td>0.27</td>
<td>0.09</td>
</tr>
<tr>
<td>H_MH</td>
<td>39</td>
<td>23</td>
<td>0.9555</td>
<td>0.0142</td>
<td>−0.805</td>
<td>0.05</td>
<td>0.56</td>
<td>0.28</td>
<td>0.11</td>
</tr>
<tr>
<td>H_LIP</td>
<td>38</td>
<td>25</td>
<td>0.9730</td>
<td>0.0160</td>
<td>−0.578</td>
<td>0.05</td>
<td>0.53</td>
<td>0.32</td>
<td>0.10</td>
</tr>
</tbody>
</table>

*Population: C, coastalfoothills; H, highlands/upper valleys.
†H, number of observed haplotypes; Hd, haplotype diversity; π, nucleotide diversity; Td, TajimaD.

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Fig. 1. Map of the Rio Grande de Nasca drainage area in southern Peru. The geographical profiles (Upper) show the maximum extend of monsoonal summer rains down the valleys into the desert and the resulting vegetation during the EH, EIP, and LIP (Upper profile) and MH (Lower profile). The map (Lower) shows the shift of the desert margin (broken line) between more humid (EH, EIP, LIP) and more arid (MH, at present) periods.
for ancient DNA analysis (Materials and Methods). Additionally, sequence data of selected individuals have been successfully reproduced in ancient DNA laboratories at the University of Adelaide and Yale University. All sequences have been deposited in the GenBank database under accession nos. KJ489100–KJ489306.

All mt haplotypes could be assigned to one of the four Native American founding haplogroups A, B, C, or D and confirmed through the analysis of mt coding region polymorphisms. In total, 98 different mt haplotypes could be determined. Summary statistics are given in Table 1 and a detailed list of all haplotype mutation profiles is in Table S1. Haplotype comparison and temporal and spatial network analysis (Fig. S1) reveal that ~61% of the observed mt haplotypes are unique to individuals, whereas 17% were shared among individuals within a population, 12% were shared between populations of the same ecogeographical area but different time periods, and the remaining 9% were matches between the highlands and coast (Table S1). Of the latter, most haplotypes were shared between the coastal LIP population and both MH and LIP highland populations (Table S1). These three populations also shared high frequencies of haplogroup B, as commonly observed in modern Central Andean populations (36), whereas the other coastal populations (C_EH, C_EIP, C_MH) exhibited low B frequencies and high haplogroup D frequencies (Table 1). Analysis of the genetic distances (F_{ST}) based on HVR I variation (Fig. 2) showed that temporally successive coastal populations from the EH to the MH were genetically close to each other (F_{ST} = 0.025–0.024) as well as to successive MH and LIP highland populations (F_{ST} = 0.009). When comparing coast to highland, EH to MH coastal populations were genetically distant from MH and LIP highland populations (F_{ST} 0.111–0.300). In contrast, the subsequent coastal LIP population is genetically close to both highland populations but shows a significant distance (P > 0.05, after Benjamini–Hochberg correction) from the preceding MH population and the other coastal populations (Fig. 2).

Population Genetic Comparison with Other Andean Populations. We further compared the analyzed ancient populations with several ancient and modern-day indigenous Andean populations (Table S2) to evaluate the population genetic relationships in the broader context of western South America. Genetic distances based on mitochondrial HVR I sequences were calculated and visualized using multidimensional scaling (MDS). The first coordinate of the MDS broadly distinguishes modern Central Andean from the modern Southern Andean and Fuegian populations. Both highland and coastal LIP populations from Palpa showed moderate to low genetic distances (F_{ST} 0.00–0.07) compared with most ancient and modern-day populations from the Central Andes and clustered broadly in the MDS (Fig. 2B). In contrast, the ancient coastal populations (EH and MH) exhibited significant, large genetic distances compared with most modern and ancient Central Andean populations, and are similar to indigenous populations from South Andean and southernmost South American populations (Fig. 2C).

**Bayesian Serial SimCoal Analysis.** The observed genetic distance patterns (Fig. 2), haplotype sharing (Fig. S1 and Table S1), and changes in haplogroup composition suggest a genetic discontinuity between the MH and LIP periods on the coast and an apparent affinity of the coastal LIP population to both highland populations. We used Bayesian serial SimCoal (BayeSSC) analysis (37) to test whether the observed population genetic summary statistics (F_{ST}, haplotype diversity, and TajimaD) could be explained by a demographic scenario involving large-scale migrations between the ecogeographical areas in the settlement chamber as hypothesized by archaeological research. We tested several demographic models based on the observed genetic, archaeological, and paleoclimate data, and used BayeSSC to compute a wide range of values for a number of priors in the parameter space (Fig. S2 and Dataset S2). The ABC framework allowed us to explore priors for deme size in the past and dependent growth rate as well the timing of the migration and divergence events to maximize the credibility of the final results. We subsequently compared simulated values with the observed summary statistics to assess which model was the most likely, given the data. The Akaike information criterion (AIC) was used to evaluate a goodness-of-fit value of the range of models in light of the observed summary statistics. In addition, a relative likelihood estimate for each of the 10 basic models and their variations (SI Materials and
Methods and Fig. S2) given the data was calculated via Akaike weights (ω). High AIC values, and therefore a poor fit, were obtained for the null model (H0), which represented the coastal and highland populations as one large single population and assumed continuity at the coast from the EH to the LIP. The best goodness-of-fit value and highest relative likelihood value (68%) were retrieved for model H5a (Table 2 and Dataset S2). This model allowed for a split of the coastal deme and the highland demes before the EH and two main unidirectional gene-flow events: a moderate-to-high emigration (10%) from the coast to the highlands, followed by a massive reimmigration (25%) of people from the highlands to the coast in the Early-to-Middle LIP. Models H6–H9, which explored different proportions of migrants and alternatively in reverse directions, returned low relative likelihood values. This means that the percentage of migrants per deme and the directionality are driving the models rather than a particular ratio, which confirms the usefulness of exploring the parameter space (here priors for deme sizes and their dependent growth rates). For models involving directed migration from one deme to another, we explored the timing of the migration events by using uniform priors spanning the time between the six ancient population pools for which a gene flow between coast and highlands has been assumed (Fig. S2). The event time posteriors for model H5a with the best support were 29.3 generations before 1400 AD (∼668 AD) for the 10% migration event from the coast to the highlands, and 9.9 generations (∼1153 BC) for the 25% back-migration to the coast. The next highest relative likelihood (20%) was retrieved for H1 assuming no migration between the two demes since the population split, whereas all other alternative models again returned negligible likelihood values (Table 2). Interestingly, the two models returned a posterior value of 103 (∼1175 BC) and 102 generations, respectively, for the divergence time of the coastal and highland populations.

The overall low support for all tested scenarios assuming other migration proportions and directions confirms the directionality and timing of migration processes hypothesized in H5 in agreement with the archaeological record, and provides important estimates regarding the demographic dimension of these population dynamic processes. However, we are aware of the fact that our models represent oversimplifications and cannot consider all possible demographic scenarios. We also caution that our simulations are limited to the maternal demographic history. Archaeological and ethnohistorical evidence suggests that maternal and paternal demographic histories might have differed substantially (25). In addition, our simulations are also limited by a restricted genetic diversity due to the bottleneck during the initial peopling of the Americas. A refinement of the nuances of past demographic scenarios will be possible once nuclear genome data from past populations (especially before the MH) become available.

Synthesis of Paleogenetic, Archaeological, and Ecological Data. The Nasca developed a complex social and demographic structure in the RGND (20, 38). The increase in settlement density and the advent of a hierarchic settlement structure observed in the EIP suggest that large populations accumulated in the lower valleys in the Nasca period (12, 39), highly dependent on the availability of water and complex irrigation systems (5, 12). The observed increasing climate variability at the end of the EIP would have decreased the predictability of agricultural productivity, and subsequent extremely dry conditions in the MH lowered the hydrological discharge from the highlands, inflicting a critical tail-off in yield. The settlement density in the RGND was already low when the climate instabilities started to increase (25), resulting in abandonment of the northern RGND during the hyperarid phase starting around 750 AD (12). Model H5a suggests that around 10% of the lower-valley effective population moved into the high-altitude upper valleys and highland regions and admixed with the local populations at ∼665 AD, also marking the end of the EIP and of the Nasca culture in the northern RGND. The 10% emigration to the highlands alone cannot explain the abandonment of the research area observed in the MH, but is a positive indicator that migrations out of the northern RGND coincided with the climatic alterations. It also suggests that people might have migrated to other destinations, such as southern extensions of the Nasca region, where an increasing settlement density was reported during the EIP–MH transition (19, 39). However, based on genetic data alone, we caution that alternative demographic factors such as mass mortality cannot be excluded at this stage.

The lack of significant signs of Wari presence in most of the lower valleys in the MH (12, 38) and the substantial genetic distance between the coastal and highland MH populations (FST 0.1107; Fig. 2) suggest that the decline of the Nasca might not have resulted from invasion and replacement by the Wari, as had been suggested before (40). We propose instead that the increased fragmentation and alteration of the spatial settlement pattern of the Nasca led to a reorganization of the population to smaller organizational units and eventually to the decline of the culture. The loss of a centralized management for storage and distribution of staple goods would have increased the vulnerability of the society to crop failures. This suggests that the migration of the remaining populations of the northern RGND at the EIP–MH transition might have been triggered by the increasing climate/precipitation variability.

Our results cannot explain the first decrease of settlement density in the northern RGND beginning at ∼440 AD. The archaeological record shows slightly increased settlement activities in the northern extension of the Nasca region around the Ica Valley (25), but there is no direct evidence for larger migrations from the RGND to the north. Given that we explored the parameter space covering the proposed 440 AD event, we assume that our data will only support the strongest signal, which appears to be the emigration reaching 10% after 665 AD. In addition, previous studies showed that the coastal populations of the Nasca culture region from the Paracas Peninsula to Nasca are very homogeneous (34), which indicates that migrations from the RGND

Table 2. AIC goodness-of-fit estimates, resulting model probabilities via Akaike weights, and estimated migration event times of the demographic models analyzed with BayeSSC

<table>
<thead>
<tr>
<th>Model</th>
<th>AIC</th>
<th>Akaike weight, ω</th>
<th>Event time 1 (EIP)</th>
<th>Event time 2 (LIP)</th>
</tr>
</thead>
<tbody>
<tr>
<td>H5a</td>
<td>81.46</td>
<td>0.678197</td>
<td>29.3</td>
<td>9.3</td>
</tr>
<tr>
<td>H1</td>
<td>83.87</td>
<td>0.203759</td>
<td>n/a</td>
<td>n/a</td>
</tr>
<tr>
<td>H9b</td>
<td>85.81</td>
<td>0.077213</td>
<td>26</td>
<td>6</td>
</tr>
<tr>
<td>H5c</td>
<td>87.92</td>
<td>0.026822</td>
<td>32</td>
<td>10</td>
</tr>
<tr>
<td>H6a</td>
<td>89.82</td>
<td>0.010399</td>
<td>30</td>
<td>12</td>
</tr>
<tr>
<td>H7a</td>
<td>91.95</td>
<td>0.003575</td>
<td>47</td>
<td>6</td>
</tr>
<tr>
<td>H9a</td>
<td>102.07</td>
<td>2.276E-05</td>
<td>49</td>
<td>5</td>
</tr>
<tr>
<td>H0</td>
<td>103.31</td>
<td>1.219E-05</td>
<td>n/a</td>
<td>n/a</td>
</tr>
<tr>
<td>H5b</td>
<td>108.21</td>
<td>1.052E-06</td>
<td>37</td>
<td>6</td>
</tr>
<tr>
<td>H6b</td>
<td>111.44</td>
<td>2.100E-07</td>
<td>31</td>
<td>7</td>
</tr>
<tr>
<td>H8b</td>
<td>111.57</td>
<td>1.961E-07</td>
<td>21</td>
<td>9</td>
</tr>
<tr>
<td>H7b</td>
<td>118.50</td>
<td>6.160E-09</td>
<td>50</td>
<td>3</td>
</tr>
<tr>
<td>H4</td>
<td>125.69</td>
<td>1.686E-10</td>
<td>n/a</td>
<td>12</td>
</tr>
<tr>
<td>H8a</td>
<td>136.32</td>
<td>8.311E-13</td>
<td>36</td>
<td>4</td>
</tr>
<tr>
<td>H7c</td>
<td>141.63</td>
<td>5.824E-14</td>
<td>21</td>
<td>7</td>
</tr>
<tr>
<td>H6c</td>
<td>151.42</td>
<td>4.368E-16</td>
<td>30</td>
<td>4</td>
</tr>
<tr>
<td>H3</td>
<td>166.94</td>
<td>1.863E-19</td>
<td>n/a</td>
<td>n/a</td>
</tr>
<tr>
<td>H2</td>
<td>179.14</td>
<td>4.177E-22</td>
<td>n/a</td>
<td>n/a</td>
</tr>
</tbody>
</table>

Further result details can be found in Dataset S2. n/a, not applicable.
to the northern or southern extensions would be difficult to detect with HVR I data alone.

During the LIP, the observed low genetic distances (Fig. 2), shared haplotypes (Table 2), and identical haplogroup frequencies of populations from both ecogeographical regions (Table 1) point to a process of genetic homogenization between highland and coastal populations during this period. The best-supported model, H5a, indeed indicates an immigration of ~25% highland people to the coast around 1150 AD (Table 2). This is in line with the archaeological record showing a rapidly increasing settlement density in the lower valleys in the mid-12th century coinciding with a transient phase of increased humidity in this region and protracted droughts in the southern Peruvian highlands (5, 7, 12).

It is plausible that the enhanced hydrological conditions of the lower valleys in the LIP may have boosted the carrying capacity of the agronomy acting as a primary pull factor for this migration (5), with the increasing social and demographic complexity as a secondary pull factor (34, 41). In turn, aridity and droughts in the southern highlands in the mid-12th century and the decline of the highland empires are possible push factors. The large population size in the highlands remaining after the decline of the empires and the increasing demand of space to compensate the declining agricultural productivity would have resulted in demographic pressures. These would have forced formerly urbanized populations in the highlands to become more mobile and decrease population densities (7). Parts of the populations from the southern Peruvian highlands probably migrated to more favorable environments, namely the lower valleys of the western Andean slopes. This scenario is supported by the observed high genetic affinity of the ancient highland and LIP coastal populations from the RGND to other pre-Columbian highland and modern indigenous southern Central Andean populations (Fig. 2). However, we caution that the sample size of the ancient populations remains small and the geographical distribution is limited in the case of the coastal populations.

Depending on the type of migration, mobility can be sex-specific (41). Directed colonization and elite dominance, which presumably were used by the Wari and Tiwanaku in the MH to increase their sphere of political influence (42), mostly involve mass movements of males (41). Mass movements or folk migration are alternative forms of migration involving equal proportions of males and females (41). The fact that we find support for migration from mtDNA data, describing the female population history, therefore further corroborates our interpretation of the past population dynamic processes: Demographic pressure at the end of the EIP and in the LIP forced large numbers of people regardless of their sex to migrate.

Interestingly, ethnohistoric sources and the archeological record point to a cultural distinction of foothill and upper-valley populations in the LIP and Late Horizon in contrast to the observed genetic homogenization of both ecogeographical areas (25, 43, 44). Cultural traditions are distributed and inherited in a far more complex way than genes, and cultural borders do not reflect biological borders. Persisting strong cultural traditions in the remaining coastal populations but also the necessary economic adaptation of immigrating highland people to the different ecological conditions are only two possible explanations for the observed cultural differences but mitochondrial/matrilineral genetic identity.

**Conclusion**

Our integration of heterochronous population genetic data into the framework of an interdisciplinary dataset endorses prehistoric migration processes on the western slopes of the South Peruvian Andes. The supported migration events coincide with observed climatic alterations, which must have led to a genetic homogenization in the RGND over time. Most studies analyzing the impact of climate on the mobility and demography of ancient populations have concentrated on foraging societies. Here we reveal climatic impacts on socially complex, agricultural societies. Despite a range of anticipatory agricultural adaptations observed in the archaeological record, the instability of the coastal and highland ecosystems in southern Peru exceeded critical points in the past in which migration became inevitable to sustain the population.

**Materials and Methods**

**Samples and Ancient DNA Analysis.** Samples from new archaeological sites in the area have been obtained for this study in addition to previously analyzed samples (18, 19). The new sites included two sites from the lower valleys and the coast, Chillo (n = 25) and Monte Grande (n = 20), dating to the LIP and not represented in previous studies. Other samples from coastal/foothill sites derived from the EIP phase of Monte Grande (n = 20) and the MH phase of Los Molinos (n = 12). Additionally, 18 samples came from the MH archaeological site Tranca, located in the upper valleys (Fig. 1).

Sample preparation and DNA extractions were performed using methods similar to those in ref. 45. For each individual, we generated DNA extracts from at least two independent samples.

Mitochondrial haplogroups for the samples were determined by using a multiplex single-base extension assay described in ref. 19. To determine the mitochondrial haplotypes of the individuals, we analyzed a 388-bp fragment of the mitochondrial hypervariable region I (nucleotide position np16,021–np16,408 (cRS)) using four overlapping primer pairs as described in ref. 18.

**Bayesian Coalescent Simulations.** We used the software program BayeSSC to determine the demographic scenario that would most likely support the observed population genetic data in ancient South American populations. Our goal was to test whether the observed summary statistics could be explained by genetic drift (null hypothesis) or by population dynamic events between structured subpopulations (alternative scenarios). The 10 main demographic models (H0–H9), plus several variants of specific models (a, b, c), were tested following various hypotheses put forward by archaeologists (5, 12). BayeSSC analysis was performed in accordance with previous publications (47–49) with the following parameters of sequence evolution for all simulations: a generation time of 25 y, a fixed mutation rate of 7.5 × 10−8 substitutions per site per generation (50), a transition:transversion ratio of 0.9841 (51), and a gamma distribution of rates with shape parameters of 0.205 (theta) and 10 (kappa) (50). We applied a uniform distribution to estimate the effective population size (N_e) for a Holocene source population (5–1,000 females) in our region before the LIP (448 generations (H0) or 81–500 generations (H1–H9)) exploring the time since the peopling of South America (3), and one (panmixia model H0) or two (alternative coast versus highland models H1–H9) South American demes (1,000–1,000,000) at time 0, which is set here at approximately the late LIP (1400 AD) based on the youngest ages of our LIP samples. Migration events were set to explore the number of generations between each of our ancient pools. For further information on the analyses and models, refer to SI Materials and Methods and Fig. 52.

**Population Genetic Analysis.** We calculated population-specific pairwise genetic distance, haplotype diversity, and Tajima’s D in Arlequin version 3.5 (52) using 207 388-bp HVR I sequences (np16,021–np16,408) assigned to one of six populations (Table 1). FST values were estimated using the Kimura two-parameter model (51) using a gamma distribution with shape parameter of 0.205 (50). We used Arlequin version 3.5 to identify shared haplotypes between the analyzed populations. To visualize the patterns of haplotype sharing, we calculated temporal networks based on the sequence data using
To visualize genetic relationships, we performed multidimensional scaling based on the distance matrices using Statistica 10 (StatSoft).

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