

# Provincialization of terrestrial faunas following the end-Permian mass extinction

Christian A. Sidor<sup>a,b,1</sup>, Daril A. Vilhena<sup>a,1</sup>, Kenneth D. Angielczyk<sup>c</sup>, Adam K. Huttenlocker<sup>a,b</sup>, Sterling J. Nesbitt<sup>a,b</sup>, Brandon R. Peacock<sup>a,b</sup>, J. Sébastien Steyer<sup>d</sup>, Roger M. H. Smith<sup>e</sup>, and Linda A. Tsuji<sup>a,b</sup>

<sup>a</sup>Department of Biology and <sup>b</sup>Burke Museum, University of Washington, Seattle, WA 98195; <sup>c</sup>Department of Geology, Field Museum of Natural History, Chicago, IL 60605; <sup>d</sup>Département Histoire de la Terre, Muséum National d'Histoire Naturelle, Centre National de la Recherche Scientifique, 75005 Paris, France; and <sup>e</sup>Karoo Palaeontology, Iziko South African Museum, Cape Town 8000, South Africa

Edited by Neil H. Shubin, University of Chicago, Chicago, IL, and approved March 27, 2013 (received for review February 4, 2013)

In addition to their devastating effects on global biodiversity, mass extinctions have had a long-term influence on the history of life by eliminating dominant lineages that suppressed ecological change. Here, we test whether the end-Permian mass extinction (252.3 Ma) affected the distribution of tetrapod faunas within the southern hemisphere and apply quantitative methods to analyze four components of biogeographic structure: connectedness, clustering, range size, and endemism. For all four components, we detected increased provincialism between our Permian and Triassic datasets. In southern Pangea, a more homogeneous and broadly distributed fauna in the Late Permian (Wuchiapingian, ~257 Ma) was replaced by a provincial and biogeographically fragmented fauna by Middle Triassic times (Anisian, ~242 Ma). Importantly in the Triassic, lower latitude basins in Tanzania and Zambia included dinosaur predecessors and other archosaurs unknown elsewhere. The recognition of heterogeneous tetrapod communities in the Triassic implies that the end-Permian mass extinction afforded ecologically marginalized lineages the ecospace to diversify, and that biotic controls (i.e., evolutionary incumbency) were fundamentally reset. Archosaurs, which began diversifying in the Early Triassic, were likely beneficiaries of this ecological release and remained dominant for much of the later Mesozoic.

biogeography | complex networks | macroevolution | biotic recovery | paleoecology

Mass extinctions are thought to reshape the composition and ecological structure of communities on a scale unparalleled by background extinction (1, 2). Within the terrestrial realm, the replacement of dinosaur-dominated communities by those of mammals after the end-Cretaceous extinction is perhaps the best-known example of such wholesale faunal reshuffling. By contrast, understanding the effects of the more massive end-Permian extinction on terrestrial community structure has been hampered by the paucity of high-quality geographic data, with nearly all studies restricted to sequences from single basins in Russia (3, 4) or South Africa (5–8). Although some broad similarities have emerged, such as the progressive aridification of the two basins (3, 8) and a heightened diversity of temnospondyl amphibians in the recovery interval (9), limited geographic sampling in each hemisphere constrains the ability of paleontologists to distinguish regional patterns from those characteristic of the individual basins. Moreover, these broadly separated areas are uninformative regarding large-scale (i.e., continent-level) patterns of faunal evolution.

Here we examine tetrapod faunal composition in five fossiliferous areas in southern Pangea approximately 5 million years before and 10 million years after the end-Permian mass extinction. We analyze changes in biogeographic structure with four metrics of taxon occurrence data. First, biogeographic connectedness (BC) quantifies the proportion of taxon-locality occurrences relative to the maximum number of such occurrences possible. Thus, at its extremes, faunas or floras where species are distributed across all localities show high connectedness, whereas when each

species occurs at only a single locality, BC has a minimum value. A second, related factor of biogeographic structure is the propensity for species to have correlated geographic ranges. To measure this factor, we used a network clustering algorithm that measures the potential to compress a taxon-locality bipartite network into subunits. These measures have several advantages over most current biogeographic methods (10, 11) because the summary measures follow directly from the occurrence data of the taxa included, rather than being filtered through a similarity measure. We also analyzed the average number of basins in which species occur as well as the proportion of endemic taxa per locality. These four measures quantify complementary aspects of large-scale biogeographic structure across the Permo-Triassic boundary.

## Data Collection

Tetrapod occurrence data were collected from the Upper Permian *Cistecephalus* Assemblage Zone (AZ) and the primarily Middle Triassic *Cynognathus* AZ of South Africa and their temporal equivalents (*SI Text*). These two zones are sufficiently separated from the end-Permian mass extinction to represent reasonable glimpses of pre- and postextinction assemblages. Indeed, the *Cynognathus* AZ is considered the first postextinction biozone to feature a stable global carbon cycle (12, 13). In addition to the Karoo, we collected data on the composition of tetrapod faunas in four other areas preserving fossiliferous Upper Permian and Middle Triassic beds, namely the (i) Luangwa Basin of Zambia, (ii) Ruhuhu Basin of Tanzania, (iii) Chiweta beds of Malawi, and (iv) Beacon Basin of Antarctica (see Fig. 1 for relative positions in southern Pangea).

Data from the literature were extensively updated by over a decade of fieldwork and firsthand research of historical collections, both of which have yielded extensively revised occurrence lists as well as a variety of new taxa (14–18) (*SI Text*). Importantly, fieldwork has been able to precisely document the stratigraphic position of fossil assemblages such that the fossiliferous Permian beds of Malawi, Tanzania, and Zambia are now considered equivalent to the *Cistecephalus* AZ in South Africa (15, 17). In addition, our occurrence data benefit from the personal study of all of the relevant fossils and thus do not rely on the literature being up-to-date. We restricted our analysis to species-level taxa, as this is the most appropriate level for biogeographic analysis. However, the current state of gorgonopsid therapsid taxonomy hampered our ability to generate a species-level taxonomic list for each region for this group, and therefore we excluded this clade from our analysis. Based on our fieldwork,

Author contributions: C.A.S. and D.A.V. designed research; C.A.S., D.A.V., K.D.A., A.K.H., S.J.N., B.R.P., J.S.S., R.M.H.S., and L.A.T. performed research; and C.A.S., D.A.V., K.D.A., A.K.H., S.J.N., B.R.P., J.S.S., R.M.H.S., and L.A.T. wrote the paper.

The authors declare no conflict of interest.

This article is a PNAS Direct Submission.

<sup>1</sup>To whom correspondence may be addressed. E-mail: casidor@uw.edu or daril@uw.edu.

This article contains supporting information online at [www.pnas.org/lookup/suppl/doi:10.1073/pnas.1302323110/-DCSupplemental](http://www.pnas.org/lookup/suppl/doi:10.1073/pnas.1302323110/-DCSupplemental).



**Fig. 1.** Paleogeographic map of southern Pangea with stars indicating the positions of the five Permian and Triassic fossiliferous areas analyzed (from left to right: Karoo Basin of South Africa; Luangwa Basin of Zambia; Chiweta beds of Malawi; Ruhuhu Basin of Tanzania; Beacon Basin of Antarctica). Corresponding faunal data from India and Namibia are discussed in the [S1 Text](#). Modern outlines of Africa (excluding Madagascar) and Antarctica (excluding East Antarctica) are highlighted. Early Triassic paleogeography (~250 Ma) is based on data originally published by Lawver et al. (40).

however, we suspect that at least two medium-sized gorgonopids are shared between the four Permian areas used here. In total, we gathered occurrence data for 62 Permian and 68 Triassic species ([Tables S1 and S2](#)).

### Network Methods for Assessing Biogeographic Structure

A number of measures have been proposed to summarize real-world networks. For example, in food webs species connectance is the number of observed relationships divided by the number that could conceivably occur (19), a property that is more generally referred to as the density of a network. To study biogeographic structure with bipartite (i.e., taxon-locality) occurrence networks, two key network statistics are required. First, to summarize the overlap of the geographic ranges of taxa, we rescale the density measure for species-locality bipartite networks. Second, to summarize network clustering (20), or the extent to which species are

grouped biogeographically, we use the higher-order relationships (i.e., taxon-taxon cooccurrences) in the occurrence network.

Biogeographic data are inherently bipartite (i.e., taxon-locality occurrences). The density of such a bipartite network can be visualized with the simple case given in Fig. 2. More formally, we rescale the bipartite density so that it is undefined when only a single locality is included (connectedness cannot be assessed for a single locality), and require all taxa to be present in at least one locality. When rescaled, the bipartite density measures the connectedness of the localities by the occurrence relationships of taxa. The rescaled bipartite density, which we introduce here as BC, is written

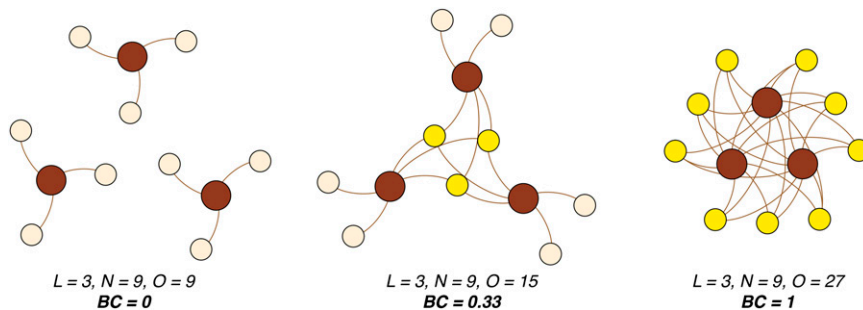
$$BC = \frac{O - N}{LN - N} \quad [1]$$

where  $O$  is the number of links in the occurrence network (number of occurrences),  $N$  is the number of taxa, and  $L$  is the number of localities. The numerator is the number of occurrences of taxa beyond a single locality (hence why  $N$  is subtracted from  $O$ ), and the denominator is the number of occurrences that could conceivably occur ( $LN$ ), minus one occurrence for each species because each species must occur at least once. This measure is bounded between 0 (when  $O = N$ ) and 1 (when  $O = LN$ ), which correspond to extreme occurrence network topologies of minimum and maximum homogeneity, respectively (Fig. 2). Between these extremes, networks with a high proportion of endemic taxa will have values closer to zero, whereas more cosmopolitan faunas will have higher values (Fig. 2). The measure does not assume that  $L$  be a specific spatial scale, or that the taxonomic rank of  $N$  be constrained to the species level.

BC is not a sufficient measure on its own to summarize biogeographic structure because a fauna or flora with different biogeographic structure but equivalent geographic range distributions can produce the same value. To more completely assess biogeographic structure, we used the map equation, an information-theoretic approach for clustering networks (20–22). When a network can be compressed into subunits with minimal information loss, it has rich cluster structure. For our purpose, a bipartite occurrence network with a low code length (measured in bits) can be compressed into more distinct biogeographic subunits than a bipartite occurrence network that clusters with a high coding length. Thus, changes in map equation scores serve as a useful proxy for differences in biogeographic clustering.

### Results

Fig. 3A shows the bipartite networks generated from our faunal occurrence data. The Late Permian network includes a substantial percentage of species (21 of 62) that are shared by at least two basins. It is worth noting that the South African Karoo has a very long history of paleontological work (23–25), which



**Fig. 2.** Bipartite networks exemplifying minimum to maximum scores of BC. Brown circles denote localities (geographic areas), tan circles denote taxa that occur at a single locality, and yellow circles denote taxa present at two or more localities. A taxon is connected to a locality if it occurs there.



conflicting dendrograms, each based on different assumptions and emphasizing various aspects of faunal similarity (10).

The network-based method we adopted here has several advantages over more traditional approaches. For example, it uses the occurrence distribution data directly, instead of filtering it through a distance calculation, and can also take into account indirect connections between communities (i.e., species that do not cooccur in a given location but are likely part of a larger regional assemblage on the basis of cooccurrences with more widely distributed species). The use of descriptive network statistics also provides more detailed insight into the underlying structure of biogeographic assemblages, and network diagrams provide an easy-to-interpret means of representing similarities and differences between faunas at various localities. In the current case, our network results show that the higher dissimilarity of Triassic communities relative to Permian ones was driven by a significant decrease in the geographic ranges of taxa and a fundamental change in the nature of the faunal connections between the basins we investigated. Interestingly, these results predict that the biostratigraphic correlation of mid-Triassic rocks should be more difficult than those of the Permian, as widespread tetrapod species became less common.

**Permo-Triassic Crisis.** The end-Permian event had a devastating effect on global biodiversity, with estimates of over 90% species extinction among marine invertebrates and over 70% in terrestrial animals (but significantly less among plants) (3–7, 26–28). Recent work has suggested that recovery was delayed because of prolonged environmental disturbance, manifested by wide fluctuations in the global carbon cycle that persisted for most of the Early Triassic (13, 29), and that substantial ecosystem diversity was not regained until about 8-million-years later in the Middle Triassic (12, 13, 30, 31). On land, the recovery of terrestrial vertebrates has been studied extensively in the Karoo Basin of South Africa. In that basin, only four therapsid genera are known to have survived the end-Permian extinction (5–7, 32), which implies that immigration substantially contributed to post-extinction faunas. Indeed, coeval rocks in Antarctica have produced fossils suggestive of a high-latitude refugium (14, 33).

The lack of well-sampled, coeval basins has hamstrung the attempts of vertebrate paleontologists to understand how the end-Permian mass extinction affected the geographic distribution of tetrapod faunas. As a result, patterns documented in the Karoo Basin have necessarily been considered representative of the southern hemisphere as a whole (5, 12, 24). Our work has shown that the four basins preserving Upper Permian fossils (Malawi, South Africa, Tanzania, Zambia) were, despite markedly different basin architecture and fill, characterized by a single, highly interconnected community during the Permian. Furthermore, the tetrapod fauna was broadly distributed (e.g., ~2,600 km between the Tanzanian and South African basins) and dominated by dicynodont herbivores, with taxa such as *Pristerodon*, *Oudenodon*, *Dicynodontoides*, and *Endothiodon* being especially wide-ranging (17).

Postextinction terrestrial vertebrate community structure changed in southern Pangea in two important ways. First, as previously shown for tetrapods (3, 4, 12, 17) the taxonomic composition of assemblages preserved in each basin underwent wholesale revision; there are no species (or genera) found in both Upper Permian and Middle Triassic rocks. Among basins, a second, larger-scale transition also occurred. Our results demonstrate the balkanization of Triassic terrestrial vertebrate faunas, with different communities present in each basin (although the depauperate Antarctic fauna is the least differentiated from that of the Karoo).

Compared with the Permian data from the same basins, the number of shared species is dramatically reduced and proportion of endemics increased. Even the Middle Triassic of Antarctica, which is by far the least well understood (14, 34), has two unique species of five total. Data for the Lower Triassic “disaster interval” are more limited, but suggest a relatively cosmopolitan tetrapod fauna with *Lystrosaurus* as its hallmark. Thus, the transition from widespread to endemic tetrapod assemblages did not directly stem from the extinction event, but instead coincided with the restoration of ecological stability during the recovery process (30, 35) (*SI Text*).

**Evolutionary Incumbency and the Rise of Archosaurs.** The timing of the archosaur radiation has received renewed attention with the recent discovery of Middle Triassic dinosauriforms (16, 18, 36) as well as Early Triassic footprints ascribed to dinosauromorphs (37) and body fossils of poposauroid pseudosuchians (38). All of these findings suggest that crown-group archosaur diversification was more intimately related to recovery from the end-Permian mass extinction than previously suspected. Importantly, this early diversification included not only phylogenetic diversification, with at least five archosaur ghost lineages drawn back to the Early Triassic, but ecological expansion into a variety of carnivorous, omnivorous, and herbivorous niches (16, 39).

Decimation of therapsid-dominated Late Permian ecosystems left vacant ecospace that was only partially filled in the earliest Triassic by *Lystrosaurus* and its associated fauna (23–25, 35). *Lystrosaurus* is broadly distributed (e.g., China, India, Russia, South Africa) and considered a disaster taxon adapted to the perturbed environments typifying most of the Early Triassic (4, 7, 33). *Proterosuchus*, an early archosauriform, cooccurs with *Lystrosaurus* and is found in the first 10 m of the Triassic in South Africa (5, 6), despite having no Permian predecessors there. Middle Triassic archosauriforms, such as *Erythrosuchus* and *Euparkeria*, are similarly found in the Karoo, but despite intensive collecting for over a century, unambiguous crown-group archosaurs have yet to be recovered from that basin. Our research in the Middle Triassic of Tanzania and Zambia has uncovered an unsuspected taxonomic and ecological diversity of archosaurs (16, 18, 36), indicating that the Karoo may not serve as a useful model system for understanding postextinction diversification. Although data rich, patterns of vertebrate recovery established within the Karoo Basin of South Africa and the south Urals of Russia fail to capture the geographic complexity of the recovery process. Data from undersampled basins show that the composition of tetrapod faunas varied to a greater degree in the Triassic than in the Permian, and that the intervening mass extinction likely dislodged therapsid incumbents and set the stage for a spatially heterogeneous recovery.

**ACKNOWLEDGMENTS.** We thank L. Nampunju, C. Sanaane, and A. Tibaijuka for facilitating fieldwork in Tanzania; A. Goulding, J. Museba, K. Mwamulowe, and S. Tolan, for support in Zambia; and the governments of Tanzania and Zambia for permission to conduct fieldwork. This research would not have been possible without access to fossils housed at a variety of museums and institutions. For access to collections, we thank the following museum personnel: C. Mehling and M. Norell of the American Museum of Natural History; B. Rubidge and B. Zipfel of the Bernard Price Institute for Palaeontological Research; M. Maisch at the Geologisch-Paläontologisches Institut Tübingen; P. Barrett, S. Chapman, and A. Milner of the Natural History Museum of the United Kingdom; T. Kemp of the Oxford University Museum of Natural History; S. Kaal of the Iziko South African Museum; and R. Asher and M. Lowe of the University of Cambridge Museum of Zoology. This research was funded by National Geographic Society Grants CRE 8571-088 (to J.S.S.) and CRE 7787-05 and 8962-11 (to C.A.S.); The Grainger Foundation (K.D.A.); the Field Museum/IDP, Inc. African Partner’s Program (K.D.A.); the Evolving Earth Foundation (S.J.N.); National Science Foundation Grants EAR 1024036, ANT 0838762, and 1146399 (to C.A.S.); and the National Research Council (R.M.H.S.).

1. Jablonski D (1986) Background and mass extinctions: The alternation of macroevolutionary regimes. *Science* 231(4734):129–133.

2. Jablonski D (2005) Mass extinctions and macroevolution. *Paleobiology* 31(Suppl 2): 192–210.

3. Benton MJ, Tverdokhlebov VP, Surkov MV (2004) Ecosystem remodelling among vertebrates at the Permian-Triassic boundary in Russia. *Nature* 432(7013):97–100.
4. Sahney S, Benton MJ (2008) Recovery from the most profound mass extinction of all time. *Proc Biol Sci* 275(1636):759–765.
5. Ward PD, et al. (2005) Abrupt and gradual extinction among Late Permian land vertebrates in the Karoo basin, South Africa. *Science* 307(5710):709–714.
6. Smith RMH, Botha J (2005) The recovery of terrestrial vertebrate diversity in the South African Karoo Basin after the end-Permian extinction. *C R Palevol* 4(6-7): 623–636.
7. Botha J, Smith RMH (2006) Rapid vertebrate recuperation in the Karoo Basin of South Africa following the end-Permian extinction. *J Afr Earth Sci* 45(4-5):502–514.
8. Smith RMH (1995) Changing fluvial environments across the Permian-Triassic boundary in the Karoo Basin, South Africa and possible causes of tetrapod extinctions. *Palaeogeogr Palaeoclimatol Palaeoecol* 117(1-2):81–104.
9. Ruta M, Benton MJ (2008) Calibrated diversity, tree topology and the mother of mass extinctions: The lesson of temnospondyls. *Palaeontology* 51(6):1261–1288.
10. Shi GR (1993) Multivariate data analysis in palaeoecology and palaeobiogeography—A review. *Palaeogeogr Palaeoclimatol Palaeoecol* 105(3-4):199–234.
11. Kreft H, Jetz W (2010) A framework for delineating biogeographical regions based on species distributions. *J Biogeogr* 37(11):2029–2053.
12. Irmis RB, Whiteside JH (2012) Delayed recovery of non-marine tetrapods after the end-Permian mass extinction tracks global carbon cycle. *Proc Biol Sci* 279(1732): 1310–1318.
13. Payne JL, et al. (2004) Large perturbations of the carbon cycle during recovery from the end-permian extinction. *Science* 305(5683):506–509.
14. Sidor CA, Damiani R, Hammer WR (2008) A new Triassic temnospondyl from Antarctica and a review of Fremouw Formation biostratigraphy. *J Vertebr Paleontol* 28(3):656–663.
15. Sidor CA, et al. (2010) Tetrapod fauna of the lowermost Usili Formation (Songea Group, Ruhuhu Basin) of southern Tanzania, with a new burnetiid record. *J Vertebr Paleontol* 30(3):696–703.
16. Nesbitt SJ, et al. (2010) Ecologically distinct dinosaurian sister group shows early diversification of Ornithodira. *Nature* 464(7285):95–98.
17. Angielczyk KD, et al. Permian and Triassic dicynodont (Therapsida: Anomodontia) faunas of the Luangwa Basin, Zambia: Taxonomic update and implications for dicynodont biogeography and biostratigraphy. *The Early Evolutionary History of Synapsida*, eds Kammerer CF, Angielczyk KD, Fröbisch J (Springer, Dordrecht), in press.
18. Peacock BR, et al. (2013) A new silesaurid from the upper Ntawere Formation of Zambia (Middle Triassic) demonstrates the rapid diversification of Silesauridae (Ave-metarsalia: Dinosauriformes). *J Vertebr Paleontol* 33, in press.
19. Dunne JA, Williams RJ, Martinez ND (2002) Food-web structure and network theory: The role of connectance and size. *Proc Natl Acad Sci USA* 99(20):12917–12922.
20. Fortunato S (2010) Community detection in graphs. *Phys Rep* 486(3-5):75–174.
21. Lancichinetti A, Fortunato S (2009) Community detection algorithms: A comparative analysis. *Phys Rev E Stat Nonlin Soft Matter Phys* 80(5 Pt 2):056117.
22. Rosvall M, Bergstrom CT (2008) Maps of random walks on complex networks reveal community structure. *Proc Natl Acad Sci USA* 105(4):1118–1123.
23. Rubidge BS (2005) Re-uniting lost continents—Fossil reptiles from the ancient Karoo and their wanderlust. *S Afr J Geol* 108(1):135–172.
24. Smith R, Rubidge B, van der Walt M (2011) Therapsid biodiversity patterns and paleoenvironments of the Karoo Basin, South Africa. *Forerunners of Mammals: Radiation, Histology, Biology*, ed Chinsamy-Turan A (Indiana Univ Press, Indianapolis), pp 30–62.
25. Hancox PJ, Rubidge BS (2001) Breakthroughs in the biodiversity, biogeography, biostratigraphy, and basin analysis of the Beaufort Group. *J Afr Earth Sci* 33(3-4): 563–577.
26. Jin YG, et al. (2000) Pattern of marine mass extinction near the Permian-Triassic boundary in South China. *Science* 289(5478):432–436.
27. King GM (1991) Terrestrial tetrapods and the end Permian event: A comparison of analyses. *Hist Biol* 5(2-4):239–255.
28. McElwain JC, Punyasena SW (2007) Mass extinction events and the plant fossil record. *Trends Ecol Evol* 22(10):548–557.
29. Shen S-Z, et al. (2011) Calibrating the end-Permian mass extinction. *Science* 334(6061): 1367–1372.
30. Roopnarine PD, Angielczyk KD, Wang SC, Hertog R (2007) Trophic network models explain instability of Early Triassic terrestrial communities. *Proc Biol Sci* 274(1622): 2077–2086.
31. Chen Z-Q, Benton MJ (2012) The timing and pattern of biotic recovery following the end-Permian mass extinction. *Nat Geosci* 5:375–383.
32. Huttenlocker AK, Sidor CA, Smith RMH (2011) A new specimen of *Promoschorhynchus* (Therapsida: Therocephalia: Akidnognathidae) from the Early Triassic of South Africa and its implications for theriodont survivorship across the Permian-Triassic boundary. *J Vertebr Paleontol* 31(2):405–421.
33. Fröbisch J, Angielczyk KD, Sidor CA (2010) The Triassic dicynodont *Kombuisia* (Synapsida, Anomodontia) from Antarctica, a refuge from the terrestrial Permian-Triassic mass extinction. *Naturwissenschaften* 97(2):187–196.
34. Hammer WR (1990) Triassic terrestrial vertebrate faunas of Antarctica. *Antarctic Paleobiology: Its role in the Reconstruction of Gondwana*, eds Taylor TN, Taylor EL (Springer, New York), pp 42–50.
35. Roopnarine PD, Angielczyk KD (2012) The evolutionary palaeoecology of species and the tragedy of the commons. *Biol Lett* 8(1):147–150.
36. Nesbitt SJ, Barrett PM, Werning S, Sidor CA, Charig AJ (2013) The oldest dinosaur? A Middle Triassic dinosauriform from Tanzania. *Biol Lett* 9(1):20120949.
37. Brusatte SL, Niedzwiedzki G, Butler RJ (2011) Footprints pull origin and diversification of dinosaur stem lineage deep into Early Triassic. *Proc Biol Sci* 278(1708):1107–1113.
38. Butler RJ, et al. (2011) The sail-backed reptile *Ctenosaurus* from the latest Early Triassic of Germany and the timing and biogeography of the early archosaur radiation. *PLoS ONE* 6(10):e25693.
39. Nesbitt SJ (2011) The early evolution of archosaurs: Relationships and the origin of major clades. *Bull Am Mus Nat Hist* 352:1–292.
40. Lawver LA, Dalziel IWD, Norton IO, Ganagan LM (2009) The PLATES 2009 atlas of plate reconstructions (750 Ma to present day), PLATES Progress Report No. 325-0509. *Univ of Texas Technical Report* 196:1–32.
41. Rubidge BS, Erwin DH, Ramezani J, Bowring SA, de Klerk WJ (2013) High-precision temporal calibration of Late Permian vertebrate biostratigraphy: U-Pb zircon constraints from the Karoo Supergroup, South Africa. *Geology*, 10.1130/G33622.1.